

# STRETTI TIDALI

*identificarli, comprenderli, rimanerne affascinati  
(criteri geologici di base per la loro identificazione nel rock record)*



Sergio.G. Longhitano, University of Basilicata, Italy



sergio.longhitano@unibas.it

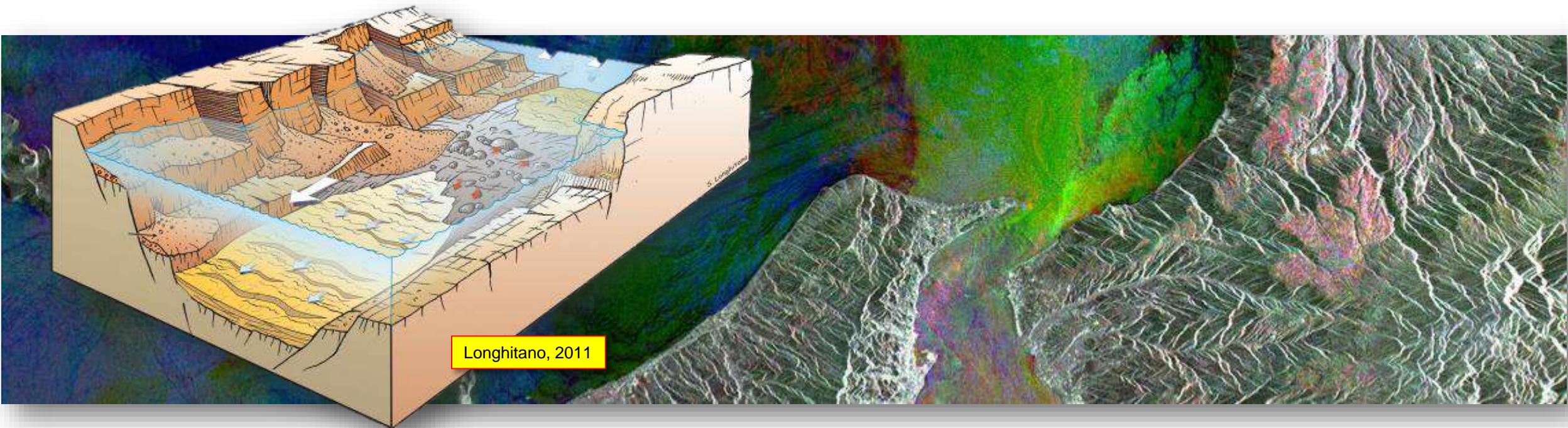
With the contribution of:

D. Chiarella, Royal Holloway University, London, UK; V.M. Rossi, CNR, Pavia, Italy; R.J. Steel & C. Olariu, Jackson School of Geosciences, Texas, USA;  
R.W. Dalrymple, Queen's University, Canada; D. Mellere, Faroe, Stavanger, Norway; A.W. Martinius, Equinor, Thrompson, Norway;  
S.A. Grundvåg, Department of Geosciences, Artic Univ. of Norway; P. Barrier, Université LaSalle Beauvais, France; M. Tropeano & L. Sabato, University of Bari, Italy; M. Pistis, University of Cagliari.



# Summary of the presentation

- Definition (geographical, geological) of a strait: why straits are important?
- Types of di straits (based on the observation of modern examples);
- Tidal Straits: Oceanographic, morphologic, bathymetric and sedimentological features;
  
- Criteria for recognising strait-fill successions in the rock record;
- What do we do not know yet about straits?





STRAITS ARE FUNDAMENTAL ELEMENTS OF MODERN GEOGRAPHY, AS THEY ARE *LOCI* AND ENGINE OF NATURAL PROCESSES AND PHENOMENA THAT TUNE AND SUSTAIN LIFE ON EARTH.

THEY ARE MORE ABUNDANT (AND RELEVANT) THAN PREVIOUSLY INFERRED!

## • Definition of STRAIT

The noun 'strait' derives from the Latin word *strictus*, meaning "to bind or draw tight", and it is commonly adopted for indicating a narrow passage of water or a tricky marine setting ([Takeoka, 1990](#)).

In Geography or Oceanography, 'strait' identifies **typically navigable waterways or shipping routes that connect two larger bodies of water that generate a restricted hydraulic cross-section, with phenomena of local water-mass convergence and related, often outstanding, sea-surface roughness and turbulences** ([Defant, 1958](#); [Pugh, 1987](#); [Pratt, 1990](#)).

Terms like '**channel**', '**pass**', '**passage**', '**sound**' or '**gateway**' are often used interchangeably with strait, although they can be referred to variable bathymetric settings or environmental conditions (e.g., [Harrison et al., 1983](#)).

Uno stretto è un braccio di acqua tra due terre che collega due bacini di acqua contigui.  
([Treccani.it – Enciclopedie on line, Istituto dell'Encyclopædia Italiana](#))

A **strait** is a naturally formed, narrow, typically navigable waterway that connects two larger bodies of water. Most commonly it is a channel of water that lies between two land masses. Some straits are not navigable, for example because they are too shallow, or because of an unnavigable reef or archipelago ([Geology Dictionary](#)).

**A strait is a narrow body of water that connects two larger bodies of water.** It may be formed by a **fracture in an isthmus**, a narrow body of land. Tectonic shifts can lead to straits like this (e.g., the Strait of Gibraltar, between the Mediterranean Sea and the Atlantic Ocean).

A strait can also be formed by a **body of water overflowing land that has subsided or has been eroded**. The Bosphorus, which links the Black Sea and the Aegean Sea, was formed this way ([National Geographic](#)).

- What tidal straits are? And why are they important?

They are:

1. influence the local oceanography, inducing **climate and sedimentary changes** in the adjoining basins;
2. prolific **fishing zones** as current exchanges also promote fish migration of great diversity;
3. advantageous and economically profitable **shipping seaways** linking marine areas separated by extensive landmasses;
4. modern analogous for subsurface many **oil-productive areas** of the world which are revealing significant reservoirs;
5. Ideal *loci* for production of **renewable energy** generated by turbines capable to produce power from the strong tidal flows;
6. preferential zones for the installation of **plumbing or wiring** for technical uses (e.g., electricity/oil/gas cabling), although the high mobility of sediment masses on the bottom of straits may often represent unstable substrates and an unwarrantable setting;
7. crucial in the **history of sedimentary basins**, as they may rapidly evolve into land passageways in case of sudden sea-level falls, having a great potential impact in the spreading of terrestrial living species or, on the contrary, inducing water masses separation and consequent marine biological diversification;
8. Important for many other reasons ...

**G O A L:**

**DEFINITION OF A DEPOSITIONAL MODEL BASED ON THE  
COMPARISON OF ANCIENT TIDALLY-DOMINATED STRAIT-FILL  
SUCCESSIONS AND MODERN TIDAL STRAITS.**

# DO WE HAVE an UNIVOQUE CLASSIFICATION for STRAITS?



Whirlpools of the Maelstrom of Saltstraumen, Nordland, Norway

# Deflection of the progradational axis and asymmetry in tidal seaway and strait deltas: insights from two outcrop case studies

SERGIO G. LONGHITANO<sup>1</sup>\* & RON J. STEEL<sup>2</sup>

<sup>1</sup>Department of Sciences, University of Basilicata, Potenza, Italy

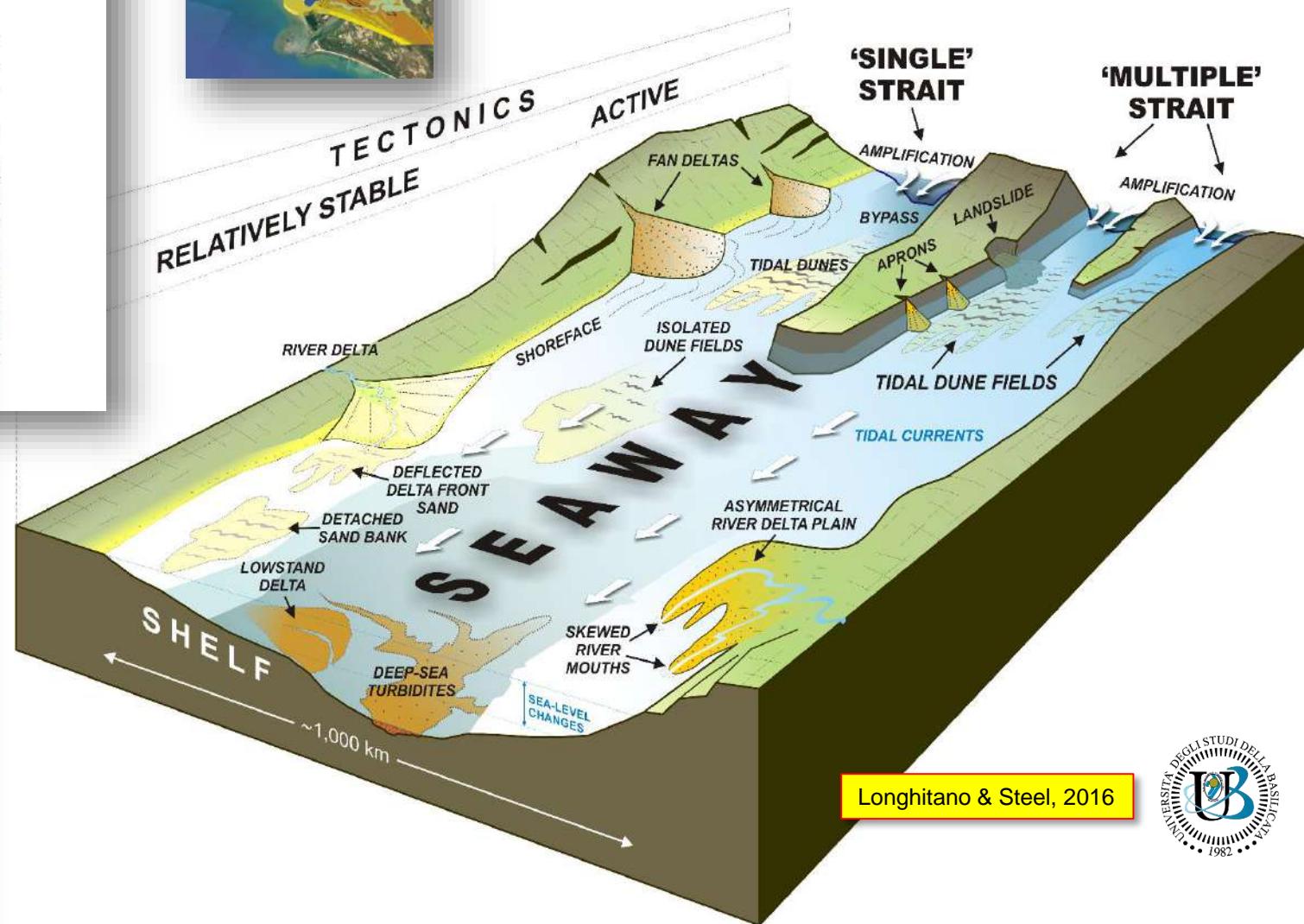
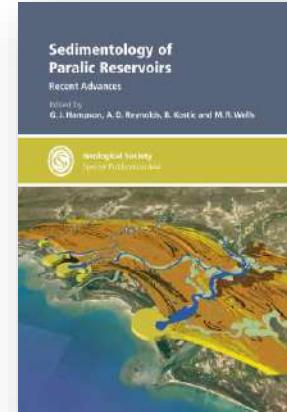
<sup>2</sup>Jackson School of Geosciences, University of Texas at Austin, Texas, USA

\*Corresponding author (e-mail: sergio.longhitano@unibas.it)

**Abstract:** Deltas represent the major sediment source for tectonically confined, tide-dominated seaways or straits. Modern examples show how along-shore tidal currents are able to modify the impinging delta shape, generating asymmetrical coastal plains, deflected delta fronts and elongate sandbanks. Seaway or strait deltas can become tide-influenced or tide-dominated, assuming physical attributes that may depart from classical models. Ancient deltas in seaways and straits can also reveal unexpected facies stacking and stratigraphies, which can be misinterpreted or attributed to basins or to tectonic settings. This study provides a synthesis of the literature on ancient and modern seaway and strait deltas in the context of the delta facies model. This study provides a synthesis of the literature on ancient and modern seaway and strait deltas in the context of the delta facies model.

## Tidal seaways and straits

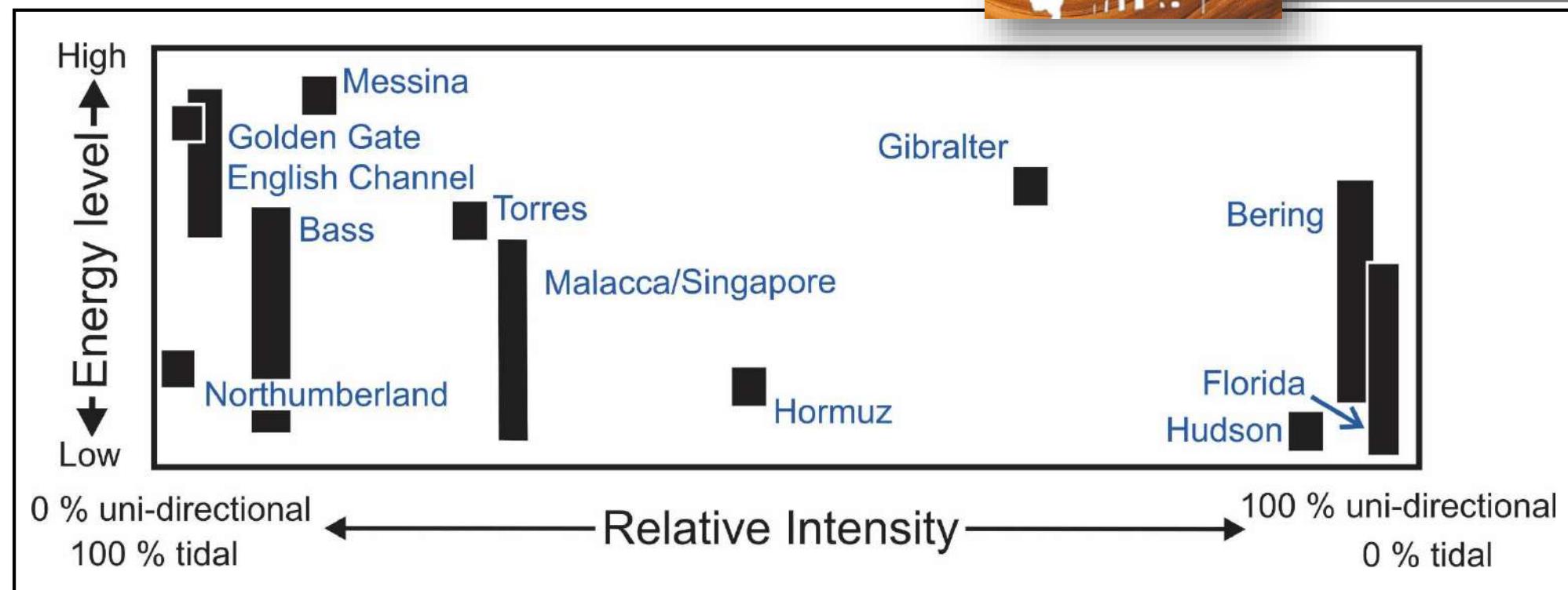
Tidal seaways are elongate marine passageways, thousands of kilometres wide, connecting two wider basins. The dominant hydrodynamic force in tidal seaways results from tidal currents flowing parallel to strike. Tidal straits are narrower than seaways and are governed by the convergence and amplification of tidal currents as a result of narrowing of the cross-sectional area of the water body (Pugh 1987). Recent studies on ancient tidal strait-fill successions in southern Italy have attempted to demonstrate that a critical cross-sectional area is the fundamental quantitative condition through which a strait becomes dominated by the amplification of a tidal flow (Longhitano *et al.* 2014). This condition, favourable to the onset of a tidal circulation, can be also matched during late transgressive stages, as has occurred recently along many of the world's coastal systems at the end of the post-Last Glacial Maximum relative sea-level rise (e.g. Longhitano *et al.* 2016).



Longhitano & Steel, 2016

- Tipi di stretti (in base a ciò che oggi si osserva)

Framework for classifying and modeling straits based on the recognition of an interplay between reversing tidal flow and some type of uni-directional current. To create a 2D parameter space that captures the variability of strait sedimentation more comprehensively, the most important variable is the **general energy level**: those with high energy are expected to accumulate abundant coarse-grained deposits, whereas those with low energy are muddy.



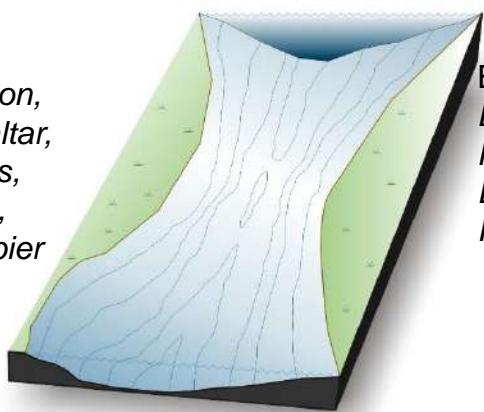
courtesy of [Dalrymple, 2019](#)

- Types of straits (based on what we observe today)

← SYMMETRIC →

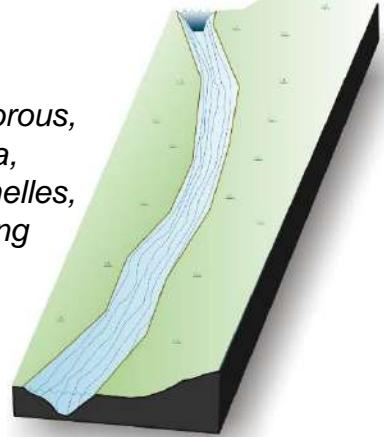
**WIDE/SEAWAY**

Es.:  
Hudson,  
Gibraltar,  
Torres,  
Cook,  
Dampier



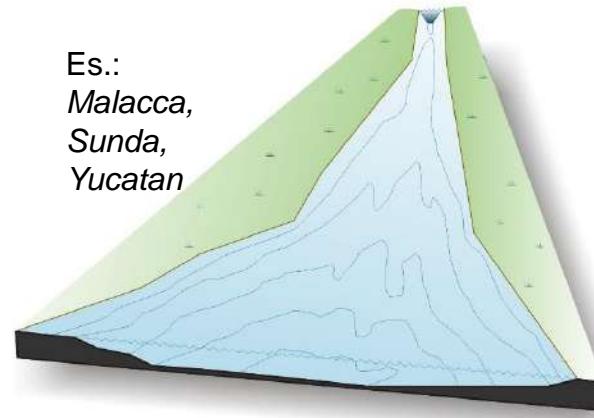
**NARROW/LINEAR**

Es.:  
Bosphorus,  
Niagara,  
Dardanelles,  
Pickering



**WIDE ESTUARINE**

Es.:  
Malacca,  
Sunda,  
Yucatan



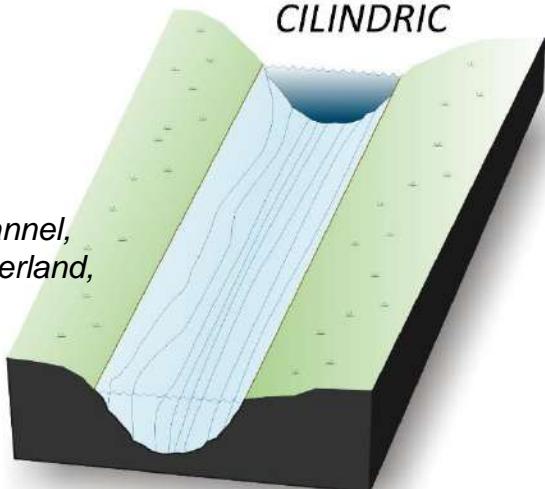
**NARROW ESTUARINE**

Es.:  
Dover,  
Øresund,  
Solent,  
Tartar



**CILINDRIC**

Es.:  
Lombok,  
Makassar,  
North Channel,  
Northumberland,  
Shelikof,  
Taiwan



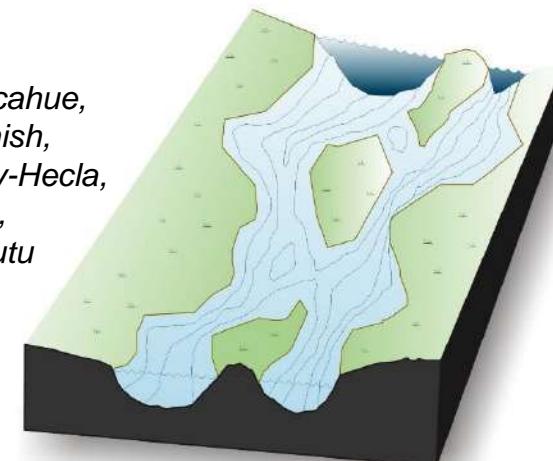
**COMPLEX**

Es.:  
Hormuz,  
Messina,  
Juan de Fuca,  
Kanmon



**MULTIPLE**

Es.:  
Dalcahue,  
Danish,  
Fury-Hecla,  
Irbe,  
Sibutu



Longhitano, unpubl.

A wide-angle photograph of a turbulent sea under a clear sky. The water is a deep blue, with numerous white-capped waves and spray. The horizon is visible in the distance.

**STRAITS DOMINATED by TIDAL CURRENTS  
(TIDAL STRAITS)**

- What tidal straits are? And why are they important?

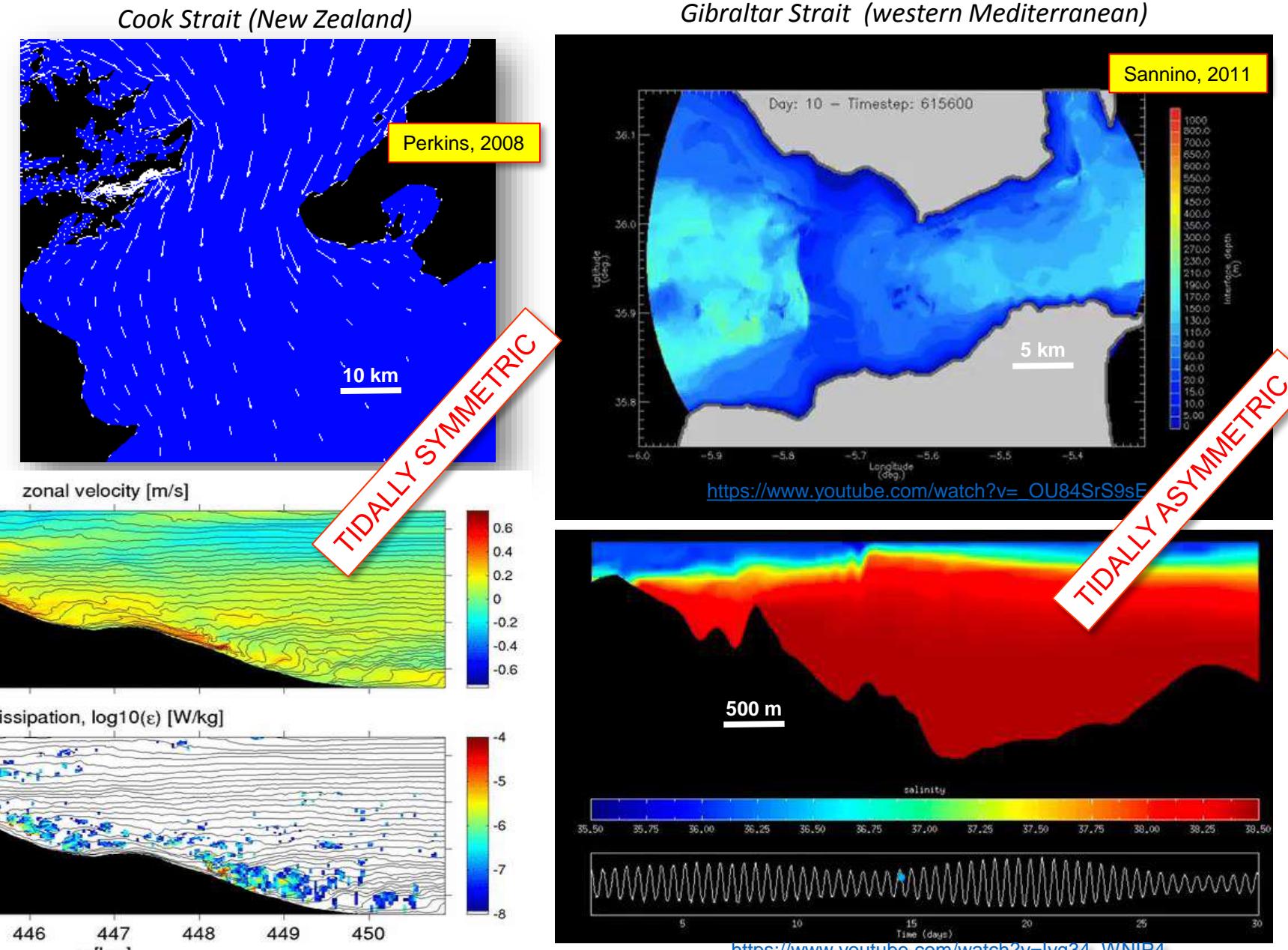
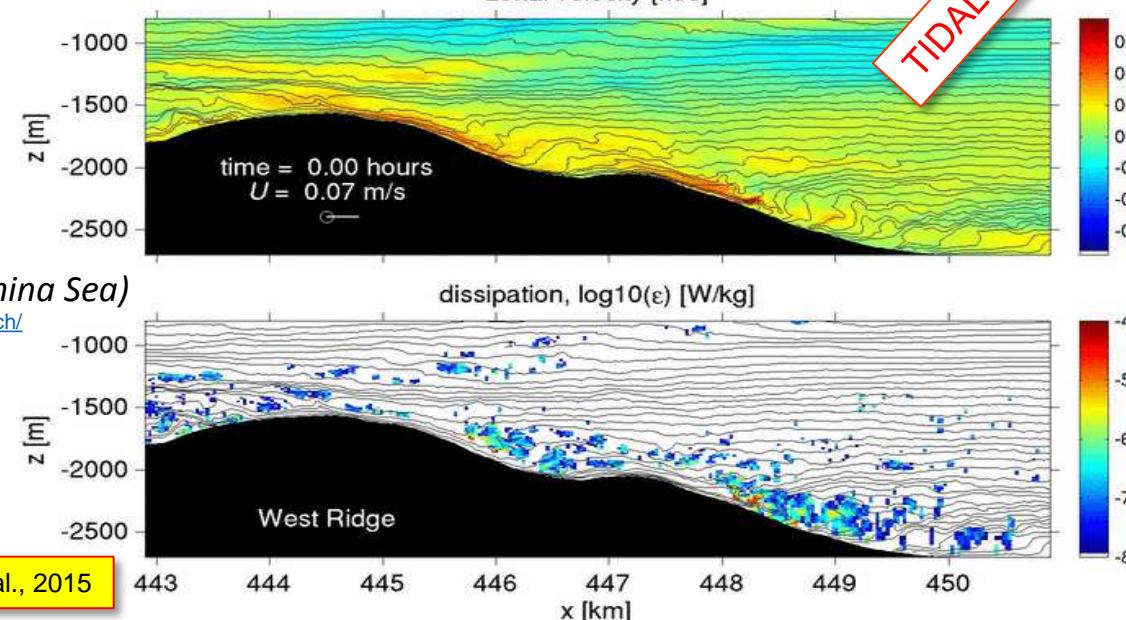
Where straits connect wider adjacent basins, whose oceanography is governed by **tidal phase opposition**, such as the modern Messina and Cook straits, the current hydrodynamics is time-dependent, responding to **a tidal cyclicity**. Tidal cycles exert a certain influence in **promoting water-mass transfers between the two interlinked basins** and a variety of associate hydrodynamics processes (e.g., turbulences, compensational flows, vortexes, etc.) (e.g., Wang, 1989; Pratt, 1990; Dalrymple, 2010).

These particular settings are thought to be characterised by **specific conditions for sediment transport and deposition**, which generally reflect a bedload divergence from a main by-pass zone, where erosion and non-deposition prevail, towards opposite zones of bedload convergence, where sediment accumulates reproducing specific bottom features (Harris et al., 1995)



- **Hydrodynamic model** (symmetric vs. asymmetric tidal systems)

The **tidal inversion** that characterizes many modern straits occurs through a change in the direction of the tidal currents. Alternations of high- vs. low-tide occurring at the same time but in the opposite, interlinked basins is the key-factor that moves large masses of water (*tidal prism*) in one direction and then in the opposite direction. This repeated cycle produces tidal currents with changing velocity/energy.



# Tidal strait w/ jetting Anisotropic LES test

Velocity (m/s)  
0.0 0.5 1 1.5 2.0

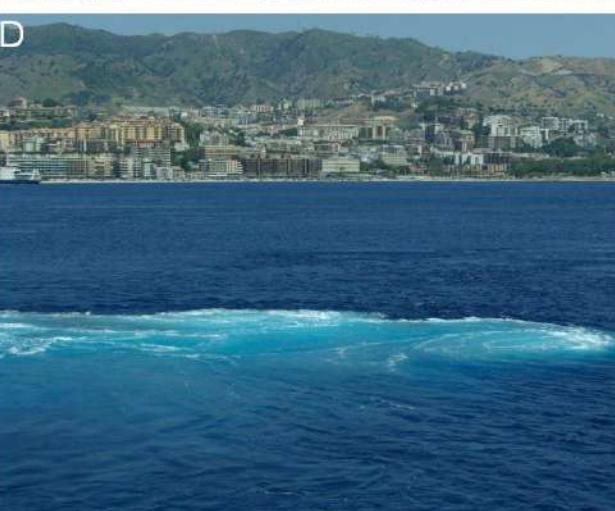
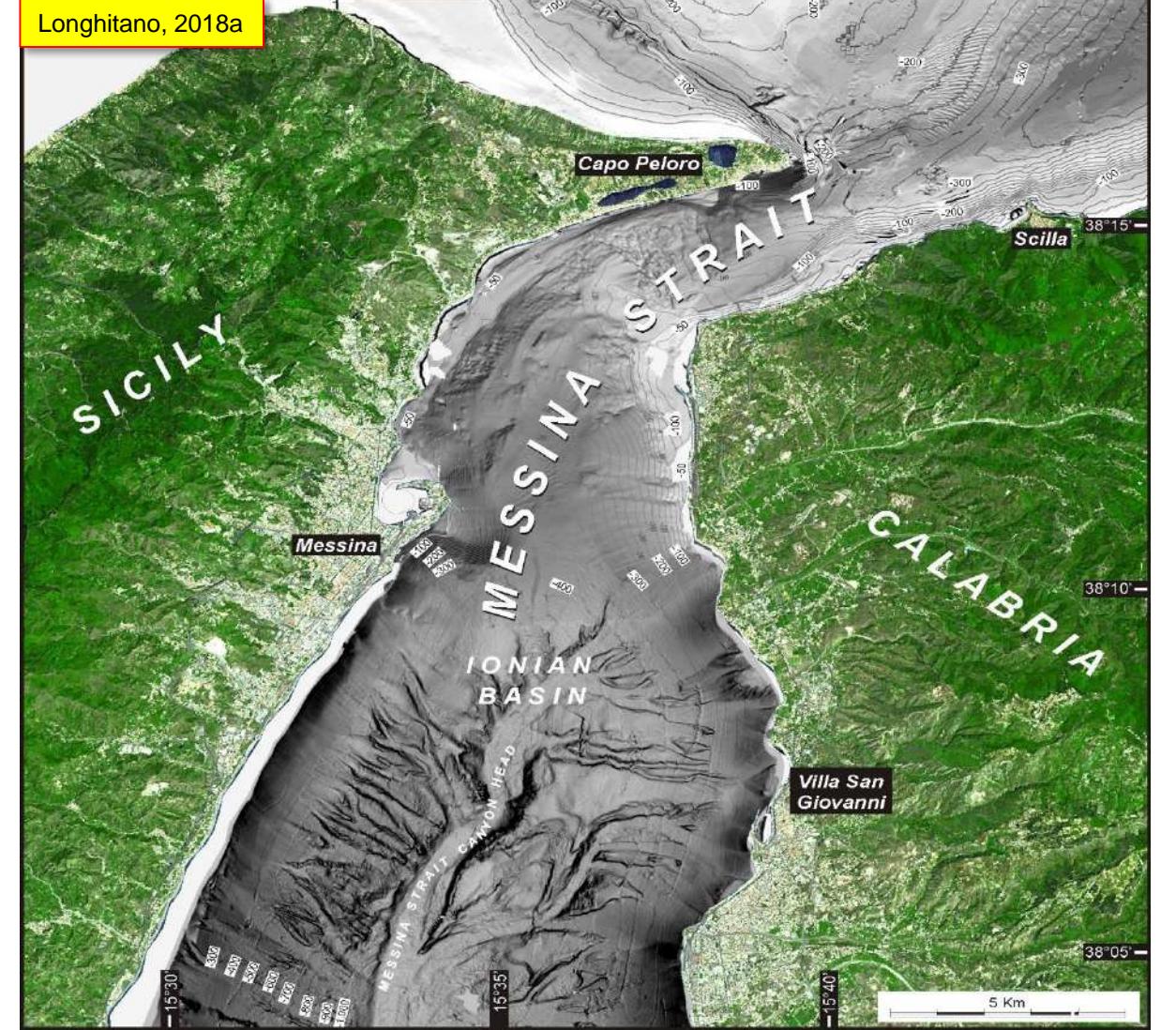


Creech, 2019

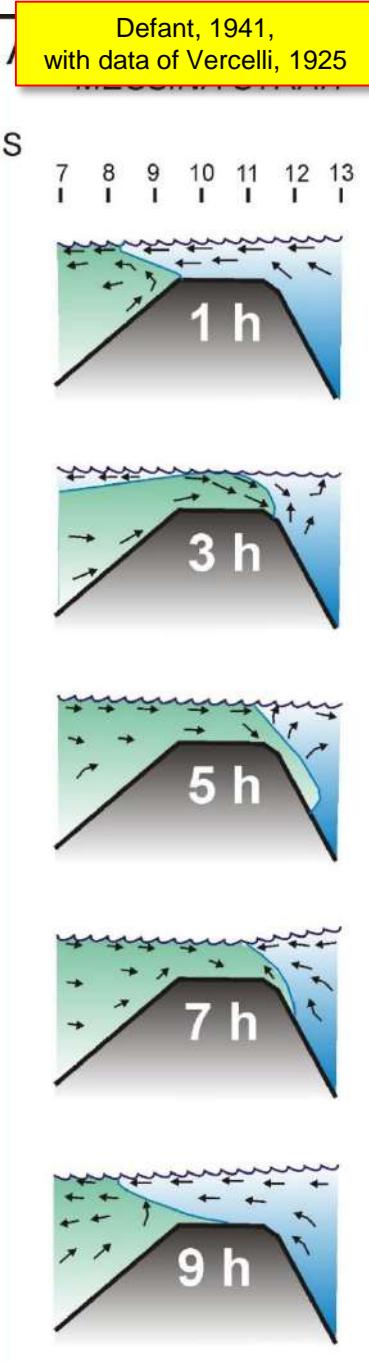
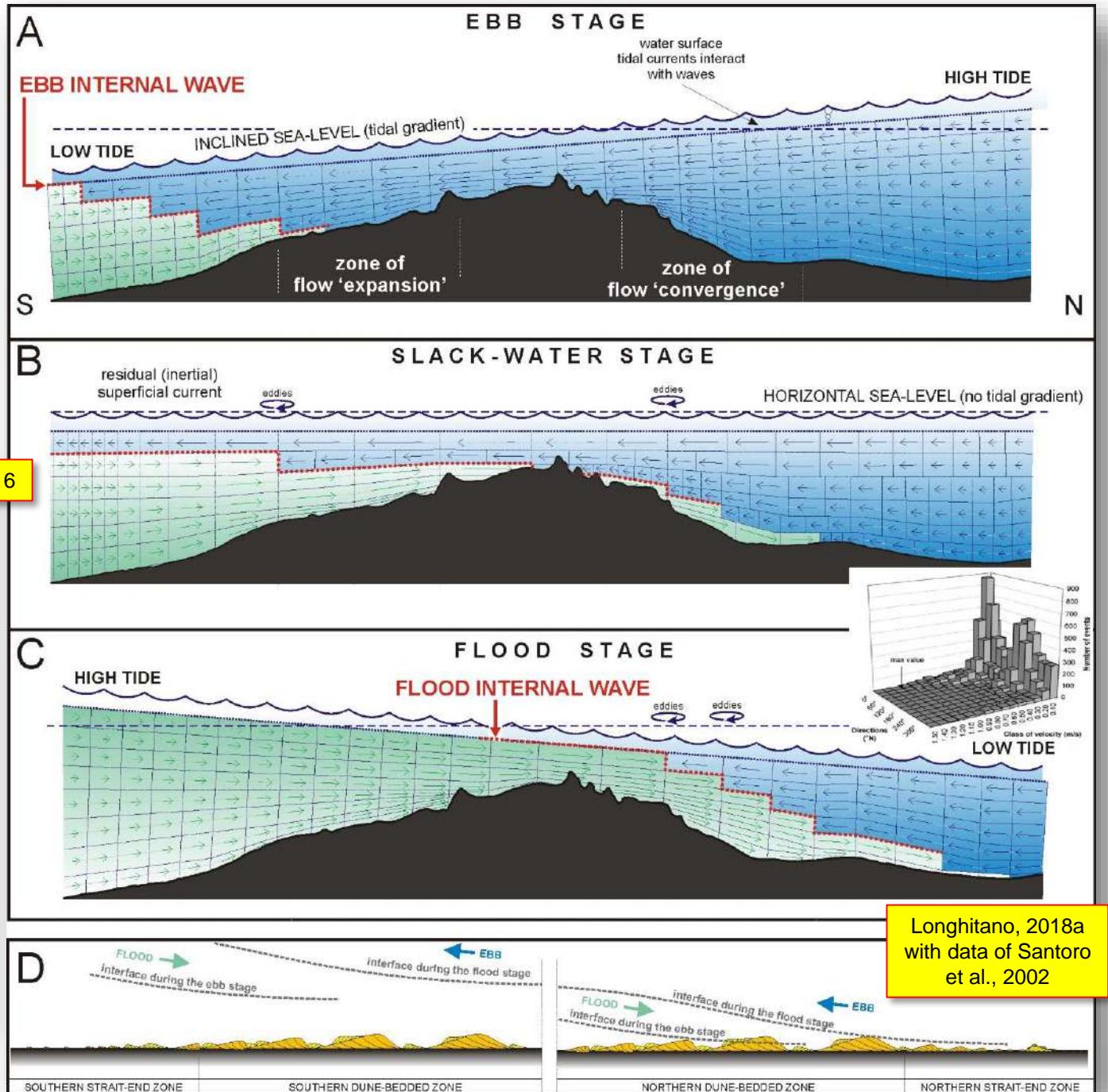
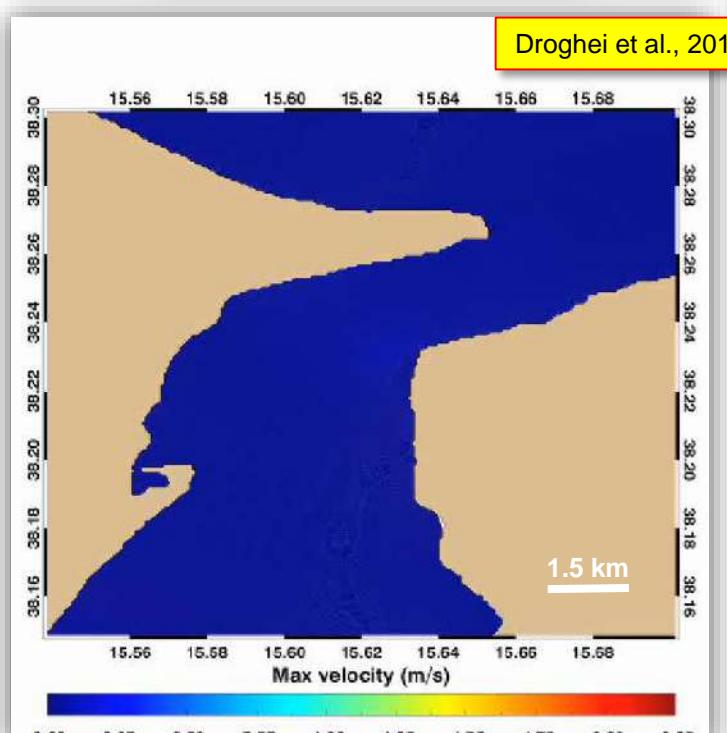
Vertical slice  
(Main channel, streamwise)

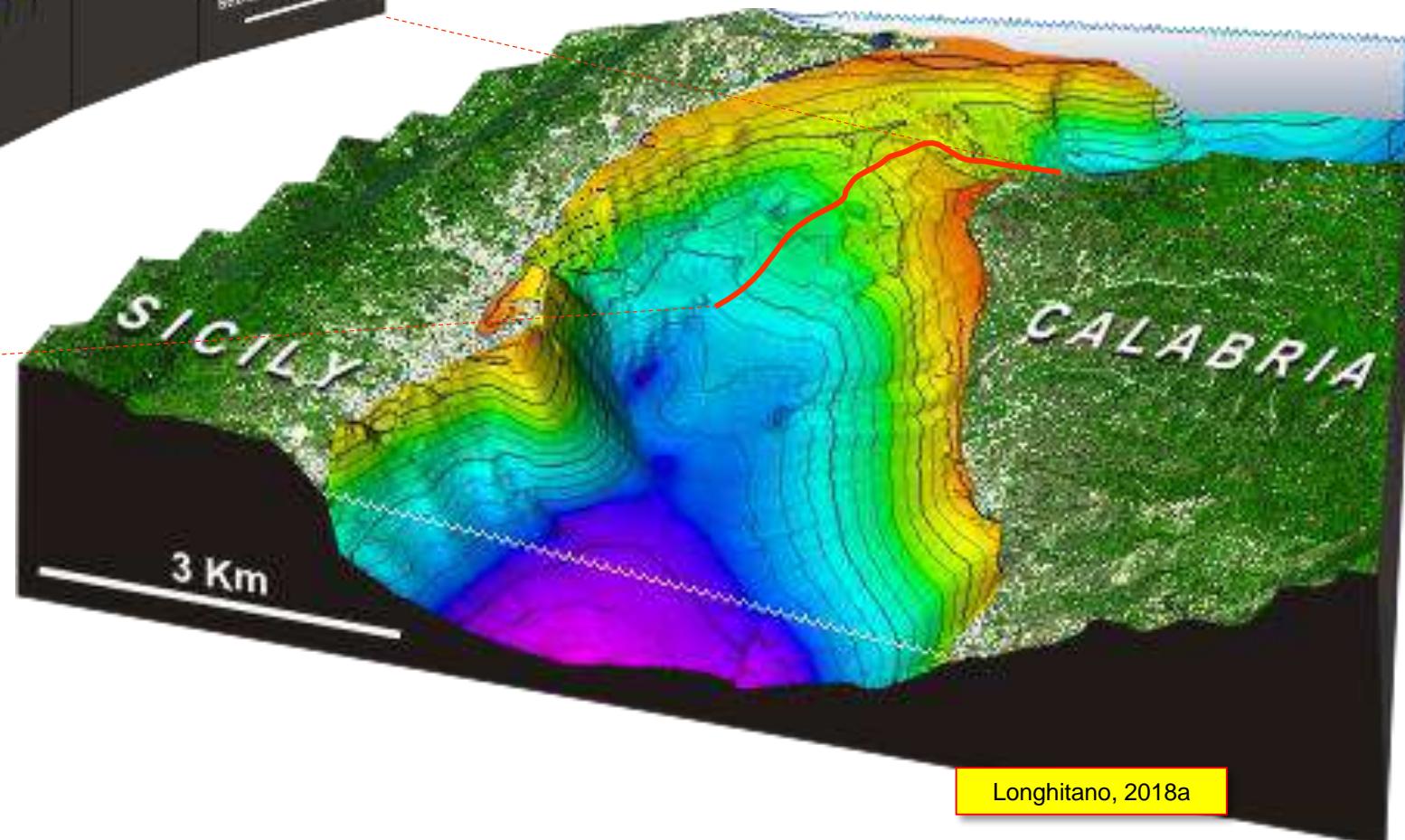
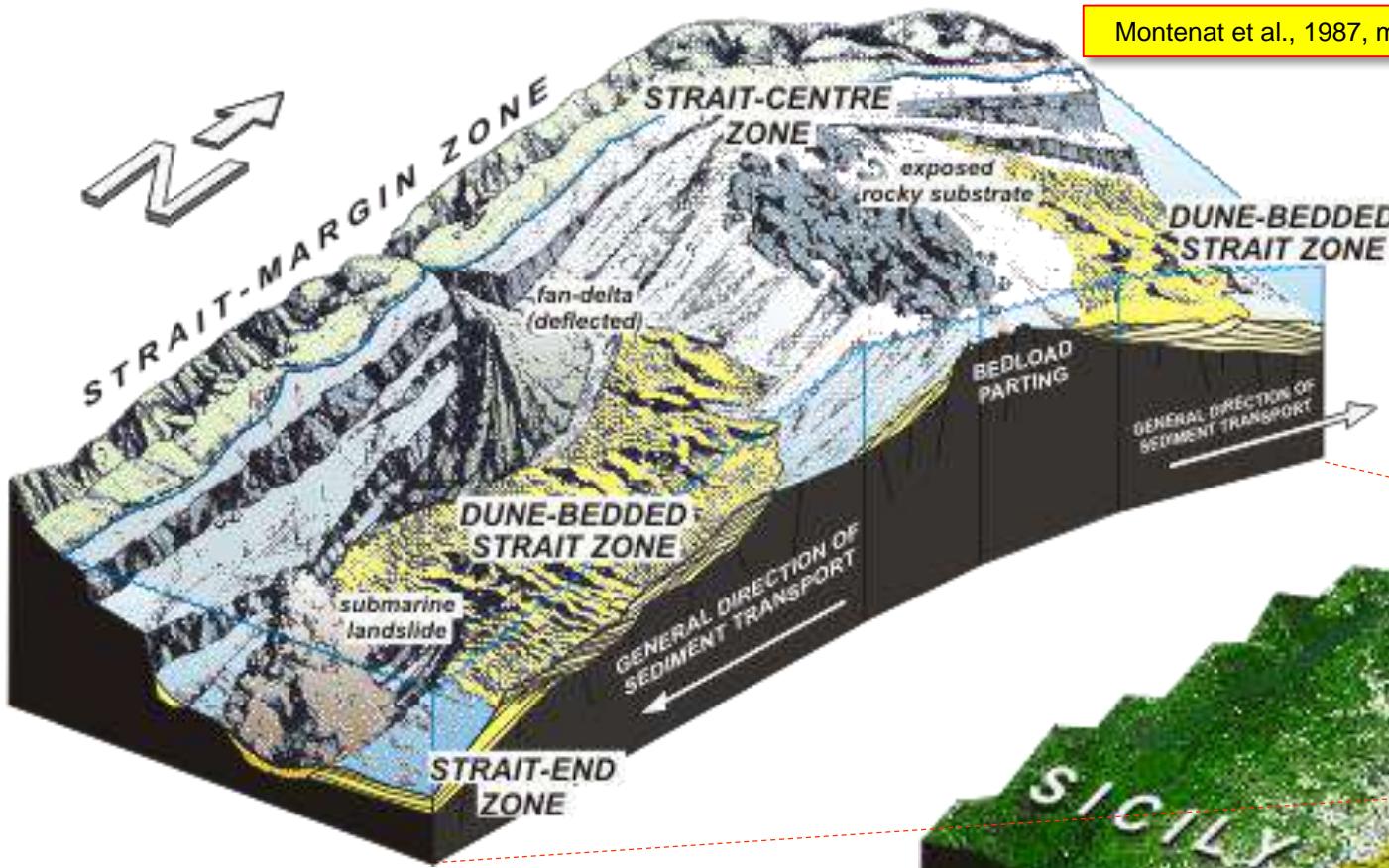
<https://www.youtube.com/watch?v=N7d2X47ltis>

Angus Creech (a.creech@ed.ac.uk)  
(c) 2019 University of Edinburgh



# Messina Strait (central Mediterranean)

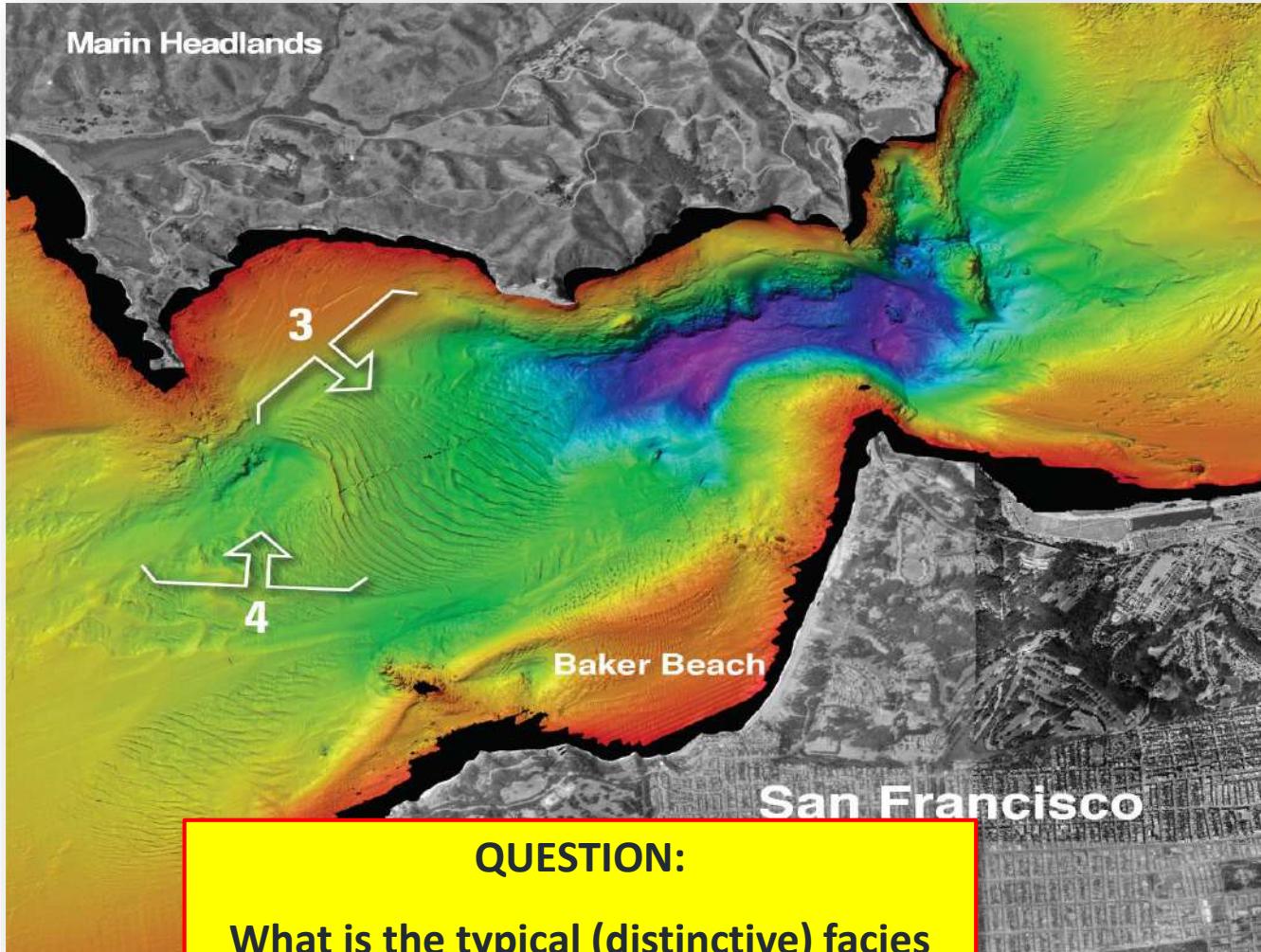




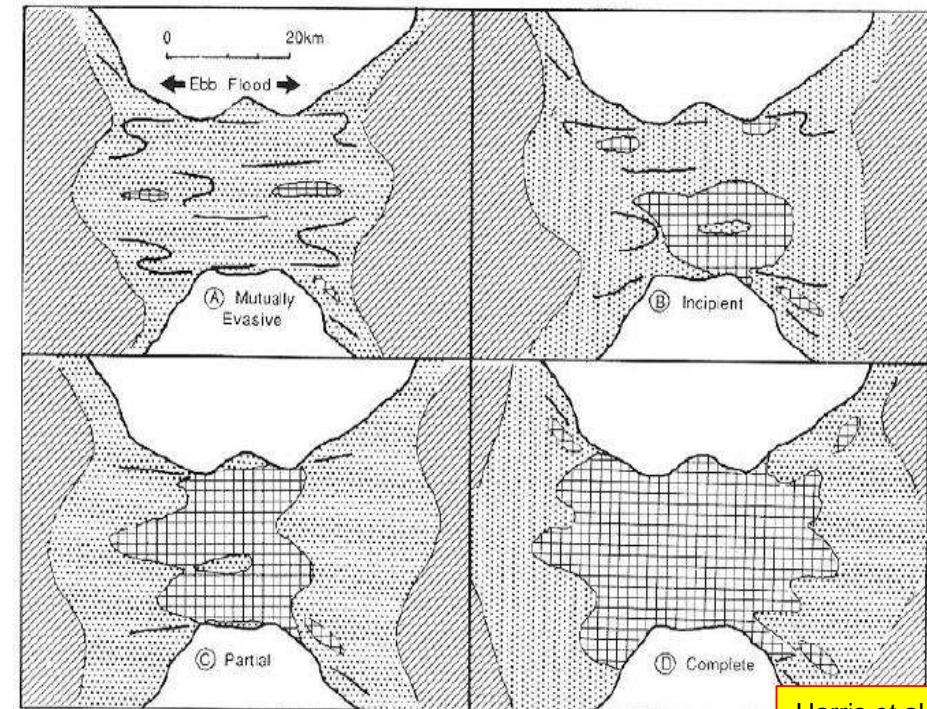
Longhitano, 2018a

- What is the sedimentary signature of tidal straits?

**SEDIMENTARY BEDFORMS** are typical features in strait bottoms. They may be **tidal sand ridges**, **sandwaves**, **dunes** and **ripples**, as result of the bed shear stress exerted by reversal tidal currents interacting with waves and other sea-water perturbations.



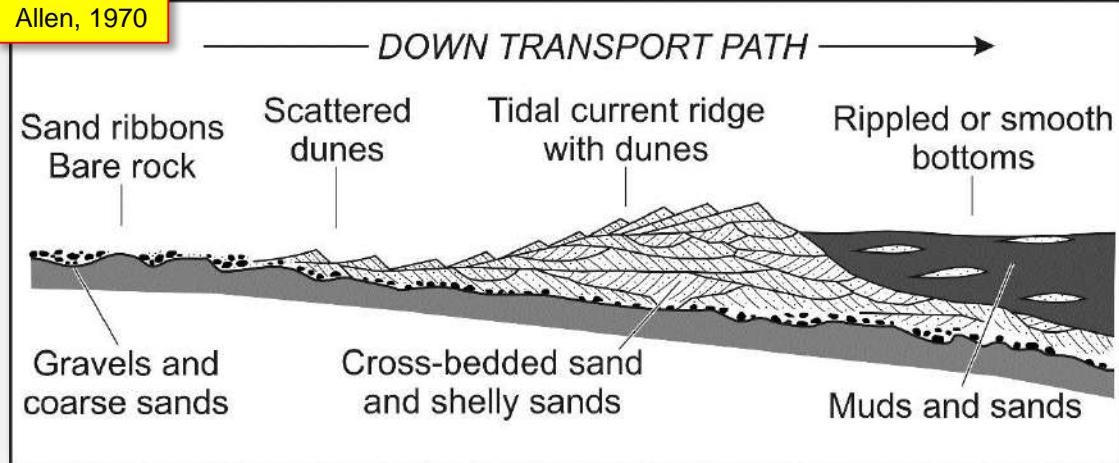
**TIDAL STRAITS** produce sediment ‘**bedload parting**’, which promotes the transport of bed material from the central (erosional) part toward either sides of the strait after flow expansion (Harris *et al.*, 1995). Bedload parting depends on the tidal phase dominance along the strait (e.g., the more symmetrical the reversal currents, the more volumetrically equivalent the depositional areas).



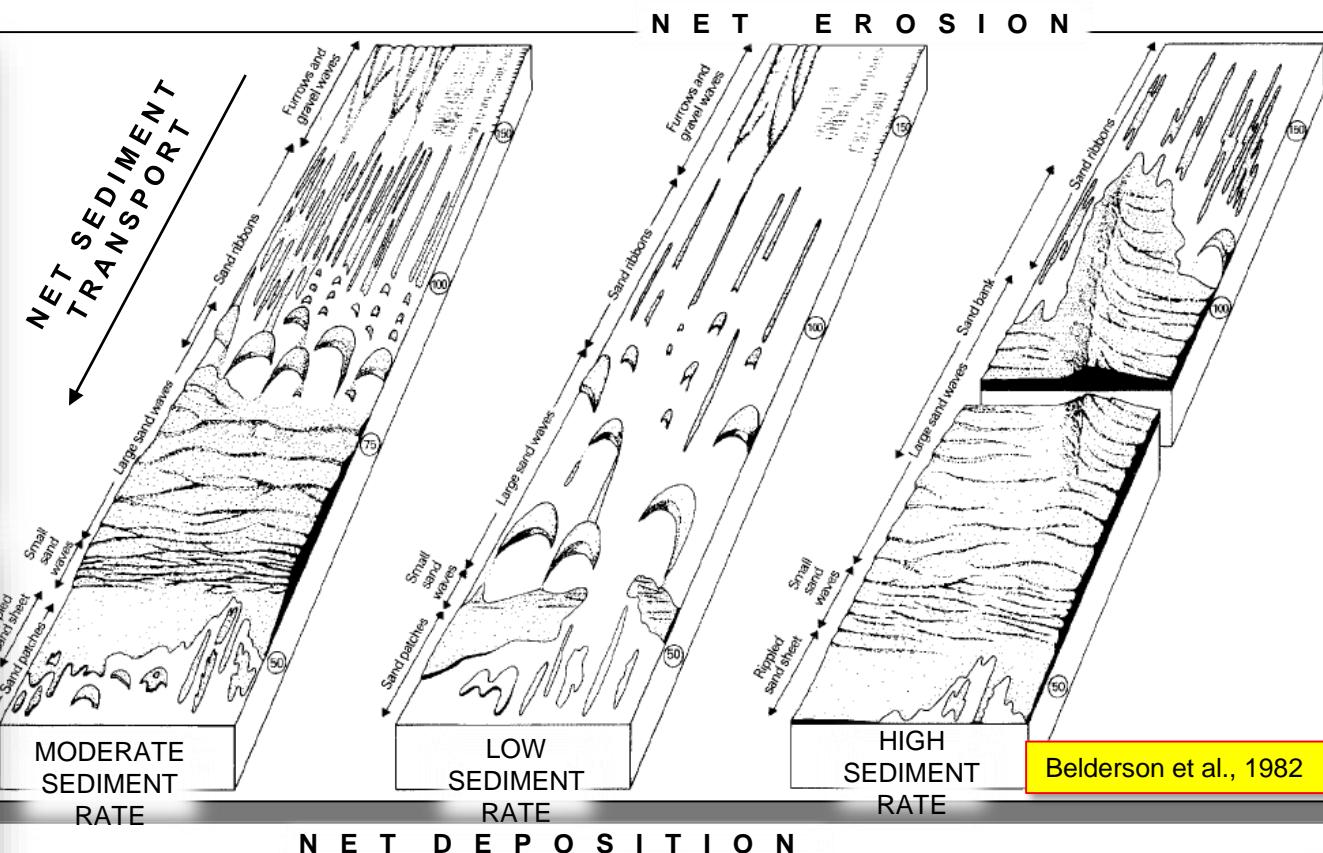
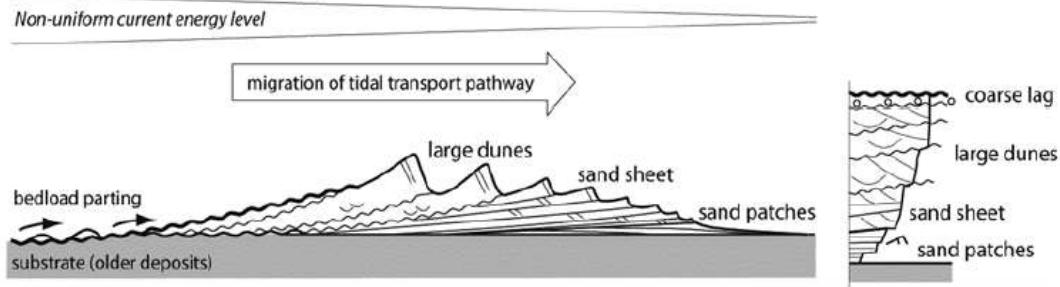
Harris *et al.*, 1995

- Previous 'embrional' models and intuitions on tidal straits ...

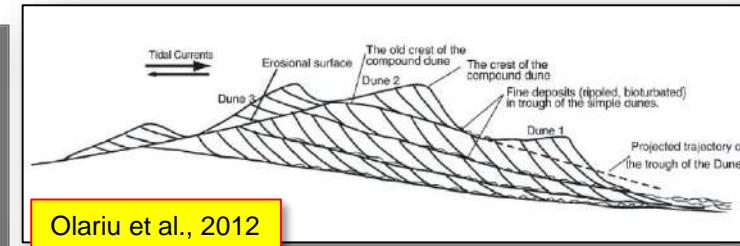
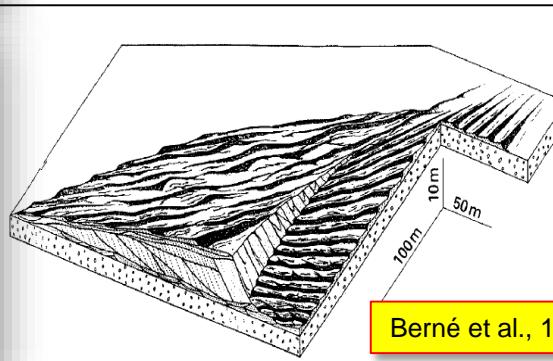
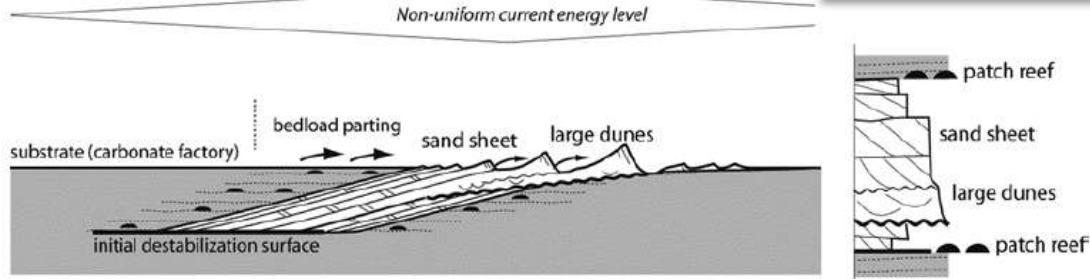
Allen, 1970



a - Deposit reworked from the substrate



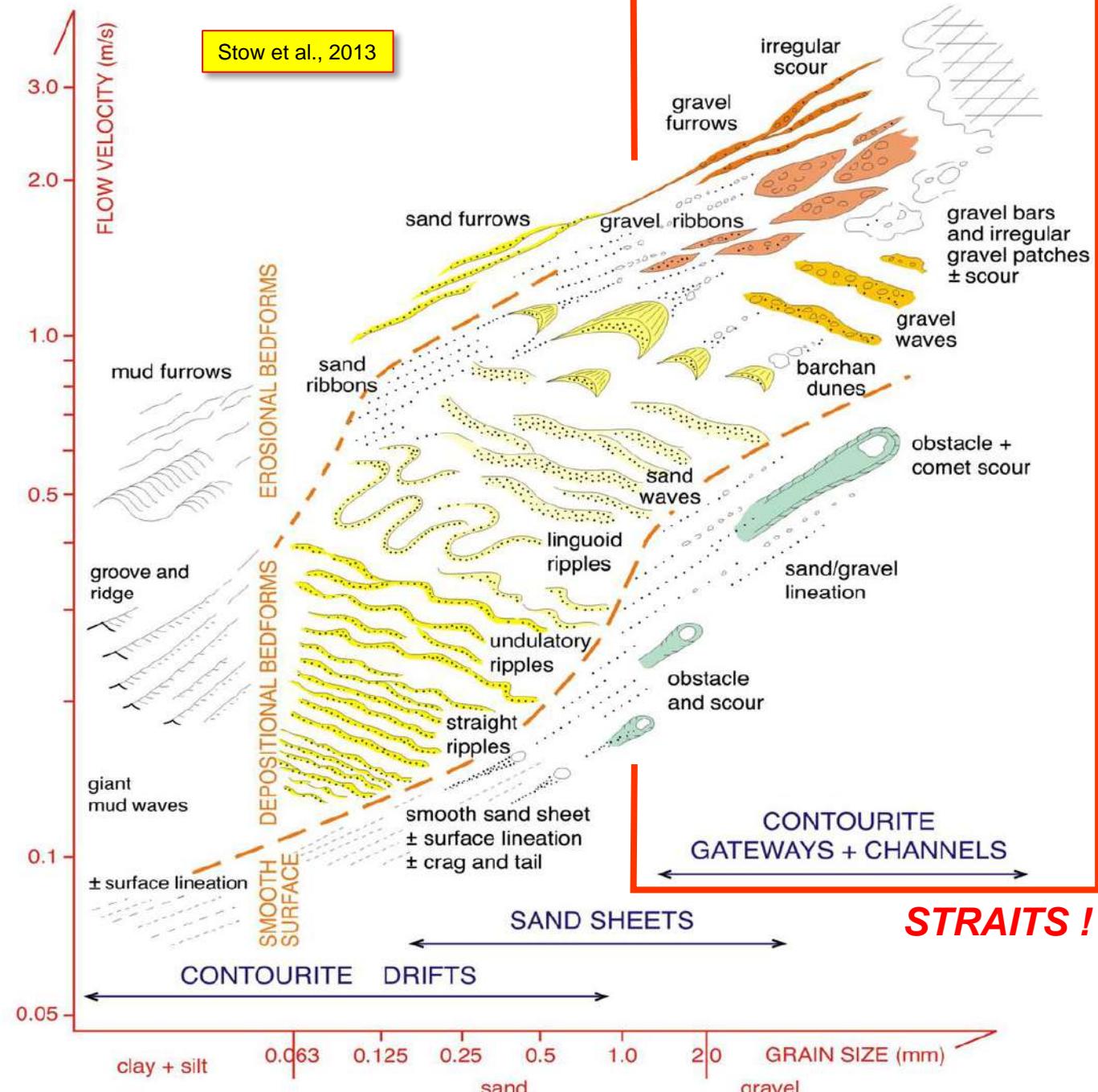
b - Deposit supplied to the substrate



Models available for modern straits have been intentionally neglected in favor of schemes, which refer to broader shelf segments. Here, typical and distinctive strait bedforms are attributed to the most proximal portions of (tidally-modulated) contouritic systems (e.g., Stow et al., 2013). This zone, here indicated as ‘gateway + channels’, perfectly **matches what we have observed in modern straits.**

However, these **bedform continua** are depth independent. In other words, the **hydrodynamic amplification** due to lateral constraining seems to be the principal driving mechanism to shape mobile bottoms into predictable bedforms.

Therefore, **tidal straits do not have preferential depth intervals to develop.**

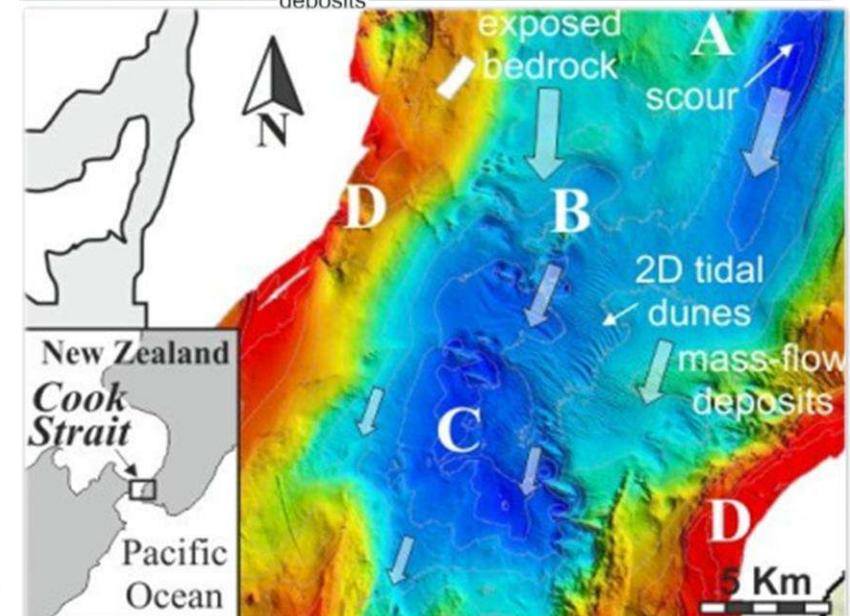
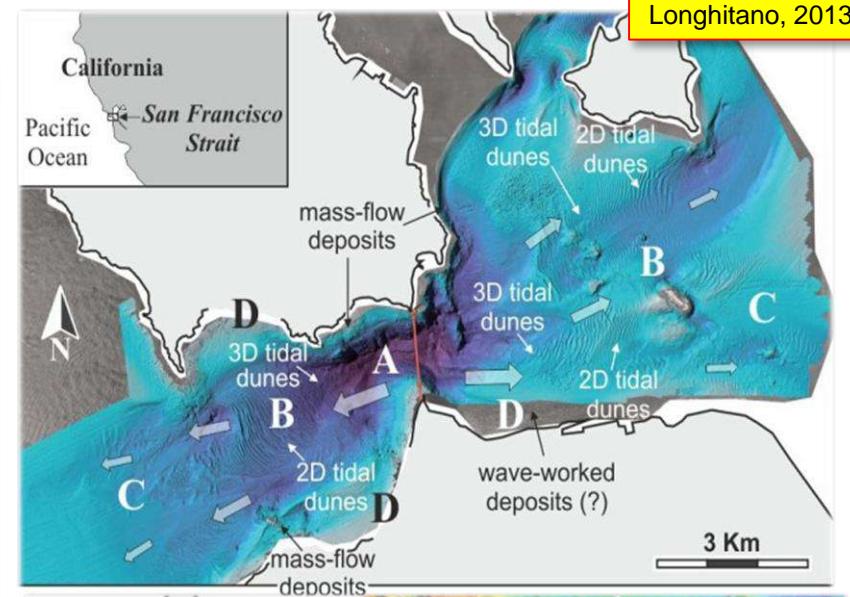
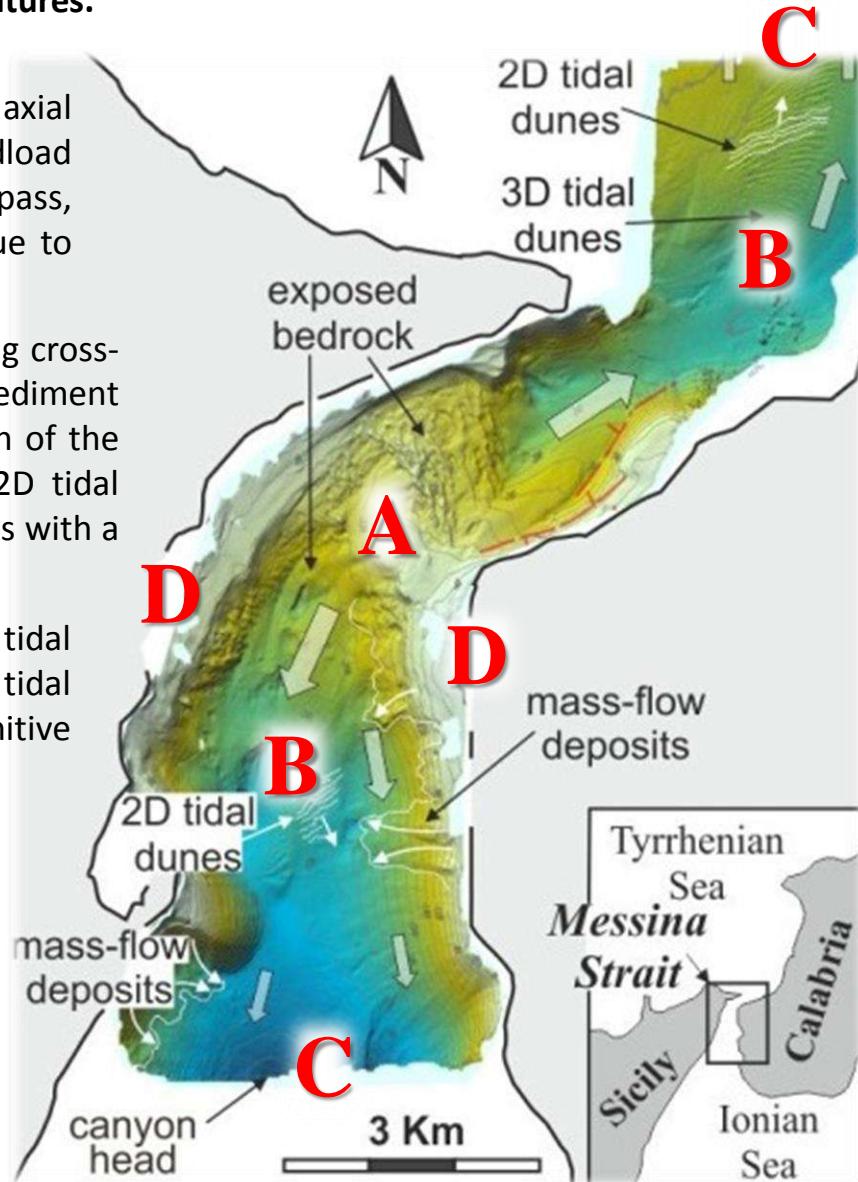


## • Depositional Zone Partitioning in Modern Tidal straits

Detailed multi-beam images of modern tidal straits allow the identification of depositional zones, each characterized by common hydrodynamics, sediment grain sizes, bedforms, and morpho-bathymetric features.

These zones are:

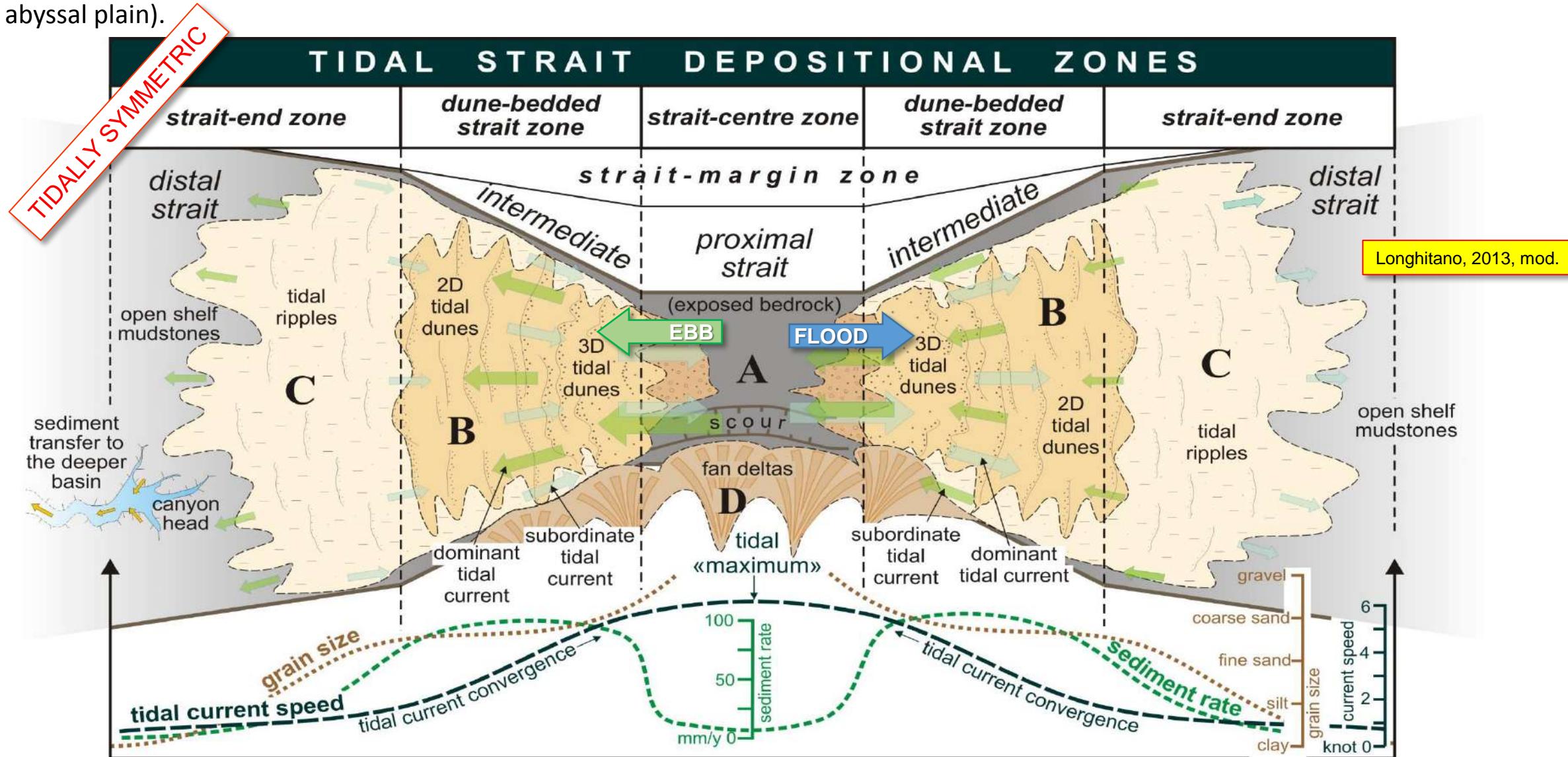
- (i) the **STRAIT-CENTER ZONE (A)**: the narrowest axial sector of a tidal strait, associated with bedload parting, tidal current maxima and sediment by-pass, with erosion or net deposition close to zero due to the highest bed shear stress;
- (ii) the **DUNE-BEDDED STRAIT ZONE (B)**: a widening cross-sectional area, is the zone of maximum sediment accumulation rate due to the initial deceleration of the tidal currents (medium to very large 3D and 2D tidal sand dunes associated with ripple-scale bedforms with a reversed or transverse direction of migration);
- (iii) the **STRAIT-END ZONE (C)**: the distal part of a tidal strait, commonly characterized by decelerating tidal currents and deposition of fines, due to the definitive enlargement of the strait cross-section;
- (iv) the **STRAIT-MARGIN ZONE (D)**: the flanks of a tidal strait, influenced by wave reworking processes in gently-sloping strait margins (e.g., the Torres Strait), or gravity-driven, sediment mass flows in straits with steeper margins (e.g., the Messina Strait). The deepest part of this zone is also influenced by tidal currents flowing along the axis of the strait.



Longhitano, 2013

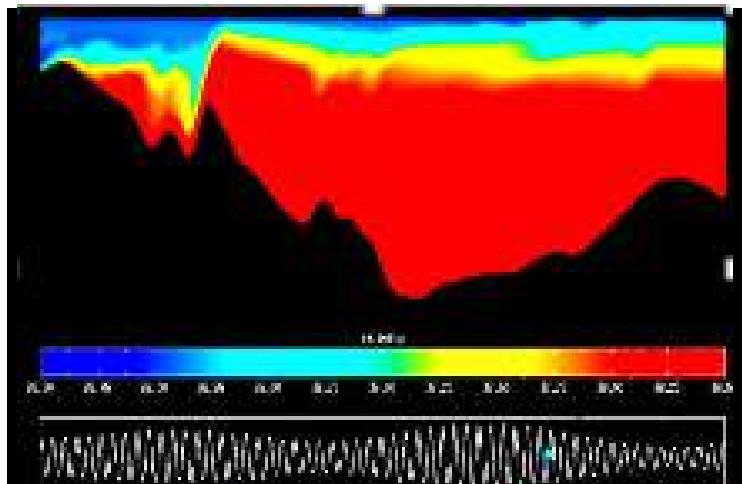
# TIDAL STRAIT: ZONE PARTITION

Modern tidal straits can be either symmetrical and asymmetrical. Examples of symmetrical systems are the Messina, San Francisco and Cook straits, where the depositional zones are equally divided by the center; the Malacca Strait is asymmetrical (its southeastern side rapidly descends into an abyssal plain).

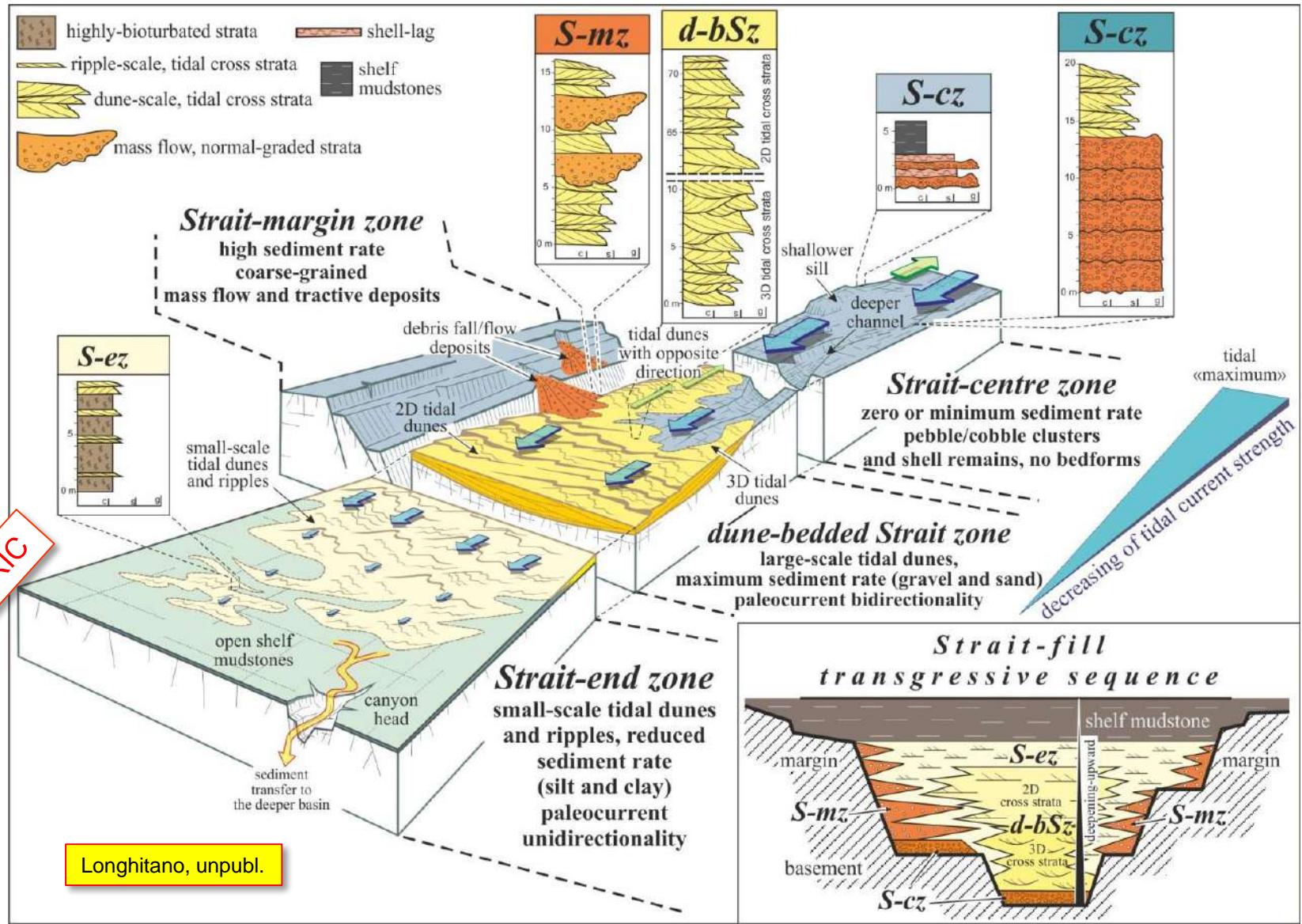


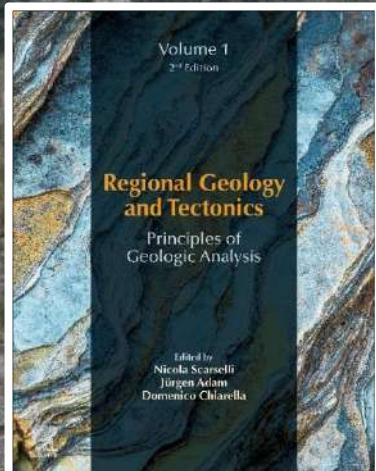
# TIDAL STRAIT: ZONE PARTITION

A symmetrical or asymmetrical configuration, which is mostly governed by the tectonic evolution of the seaway, is key in the reconstruction of ancient tidal straits, because symmetrical or asymmetrical systems generate very different volumes and distributions of dune-bedded, sand-prone deposits.



**TIDALLY ASYMMETRIC**





# CRITERIA FOR RECOGNIZING TIDAL STRAITS IN OUTCROP OR SUBSURFACE SUCCESSIONS



Siderno Strait, Calabria, southern Italy

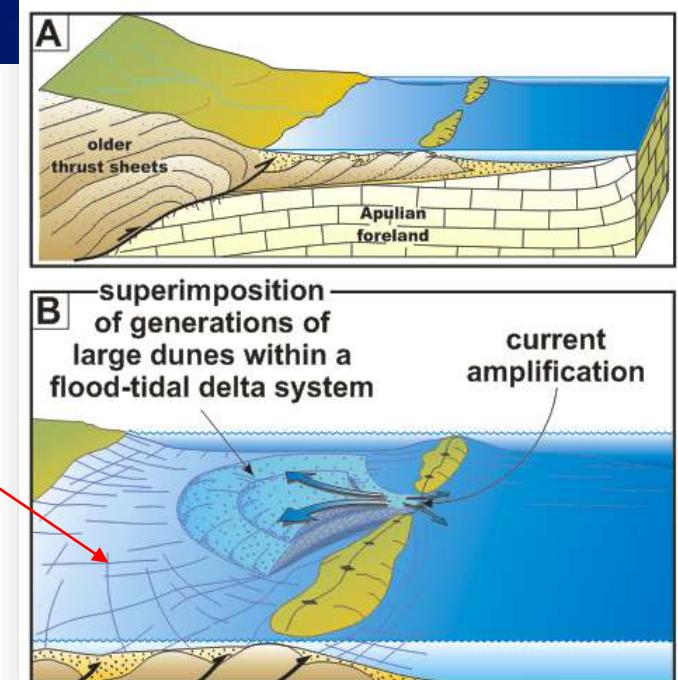
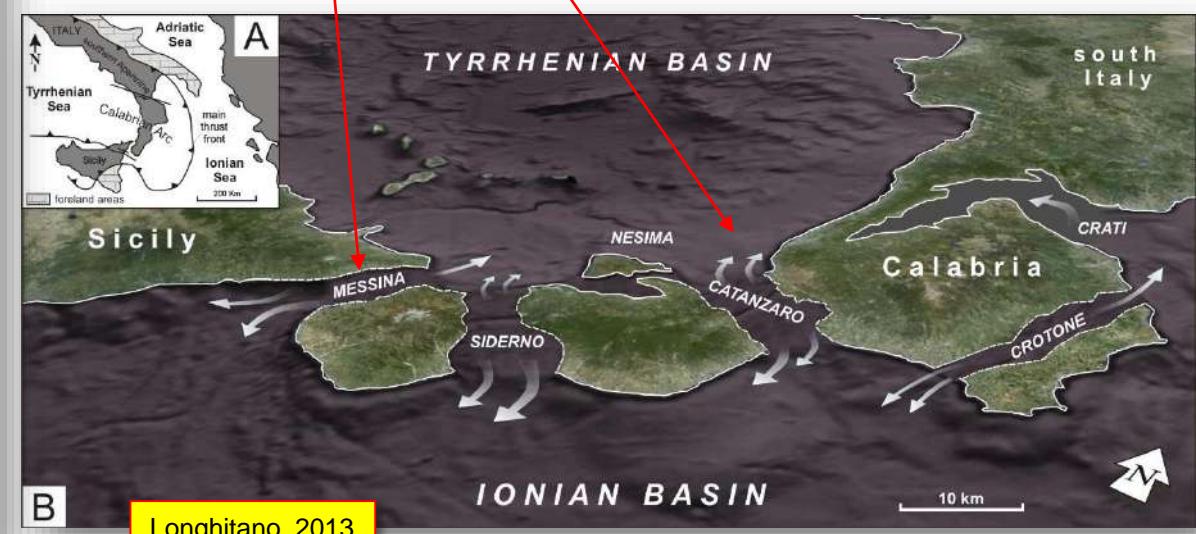
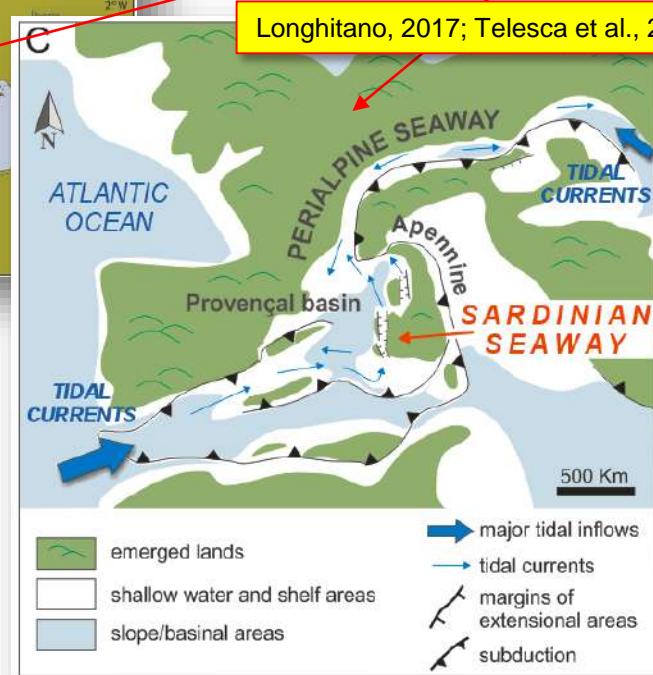
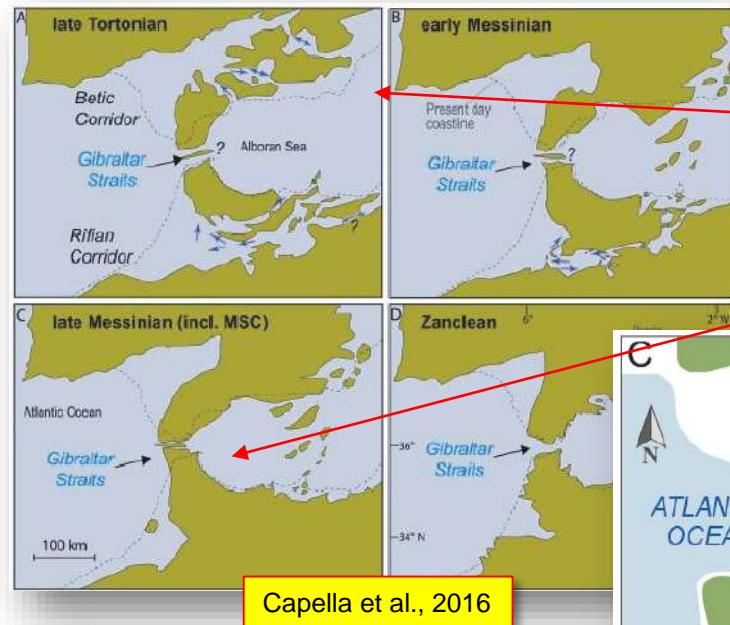
- Main areas of outcrop investigations

Stratigraphic interval: **Neogene-Quaternary**

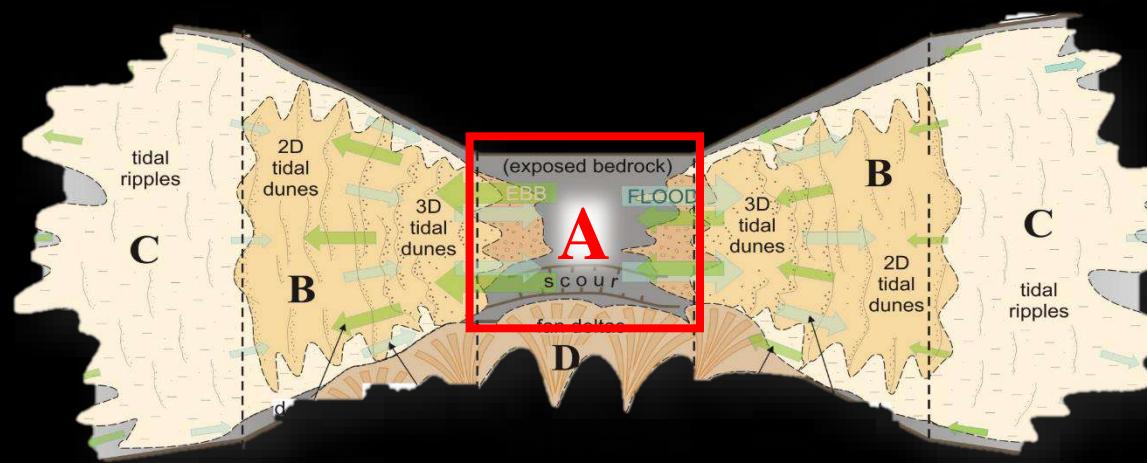
Areas: **western Mediterranean basin**

Settings: **compressional, transtensional**

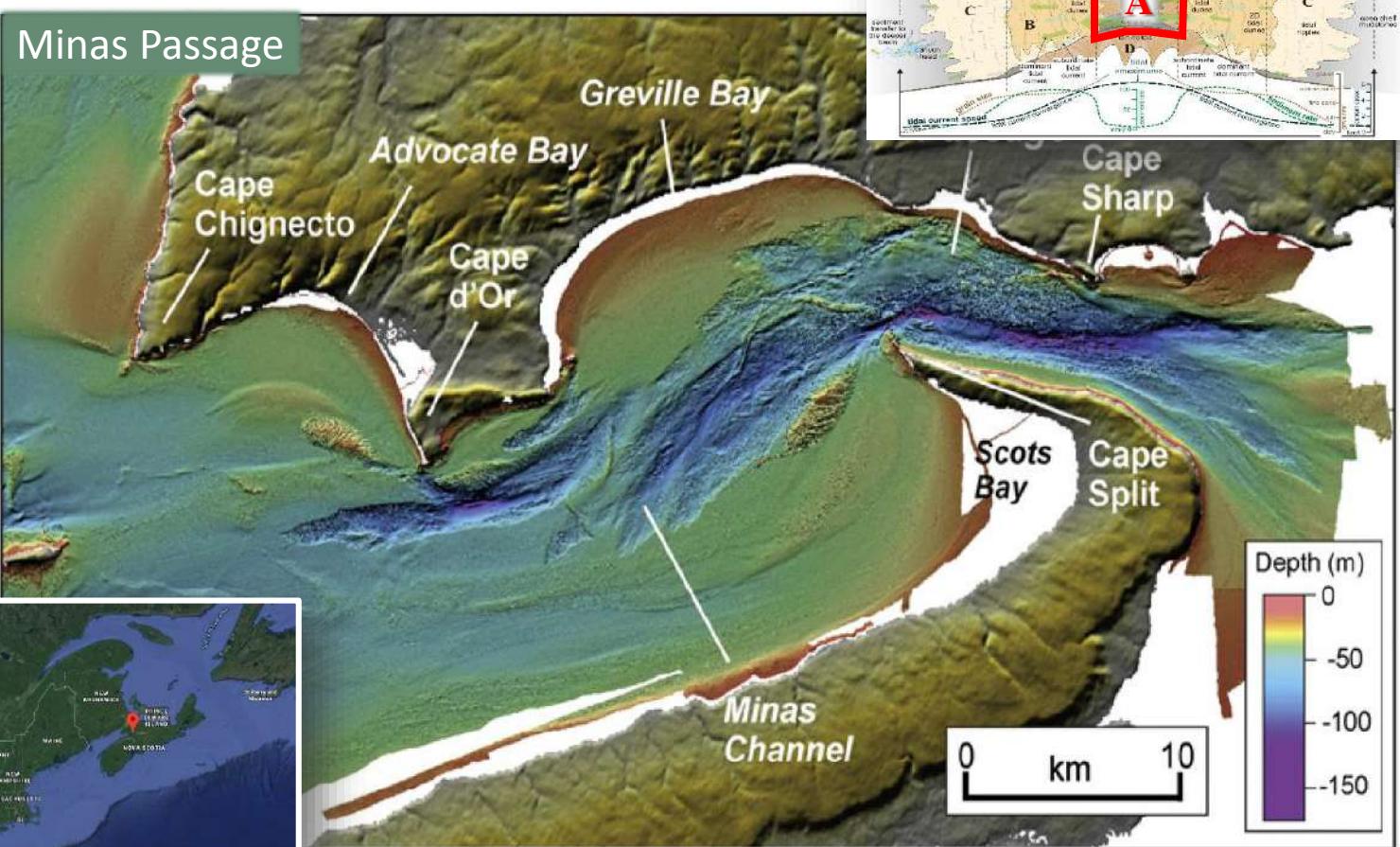
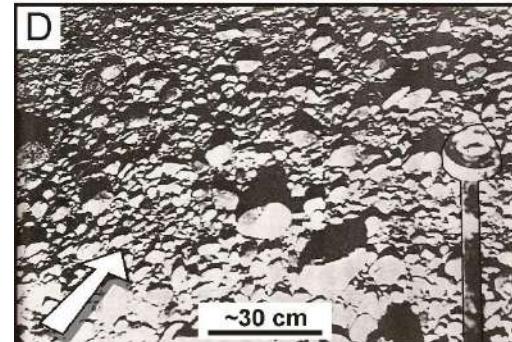
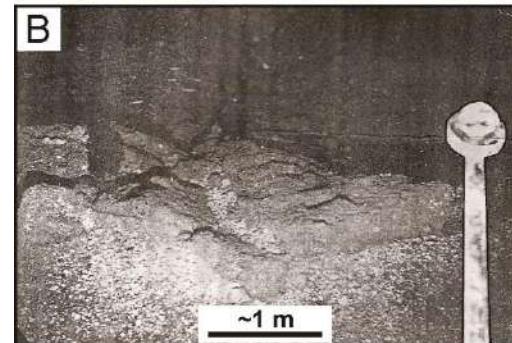
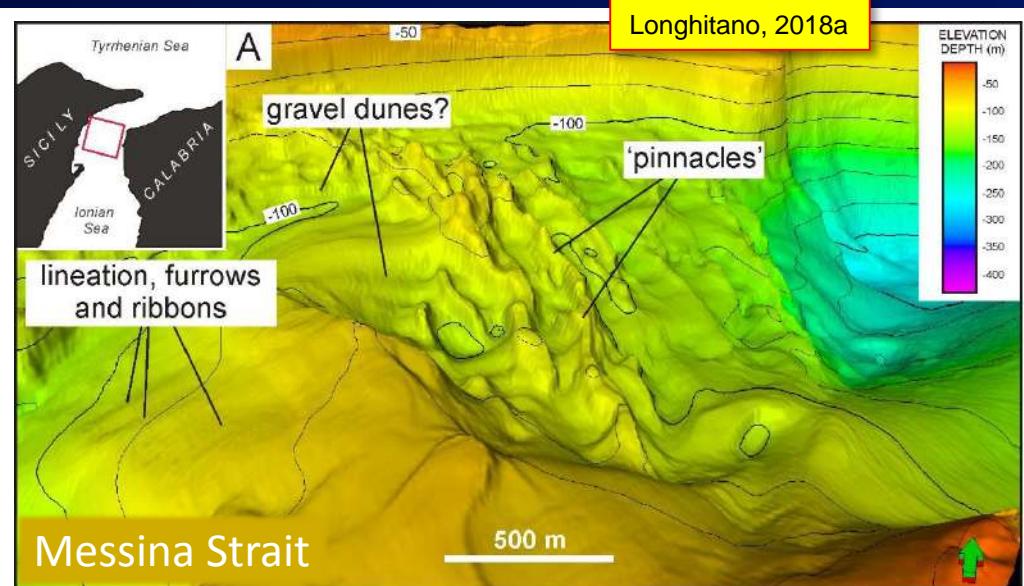
Investigated thicknesses: **25 to 200 m**



# The strait-center zone



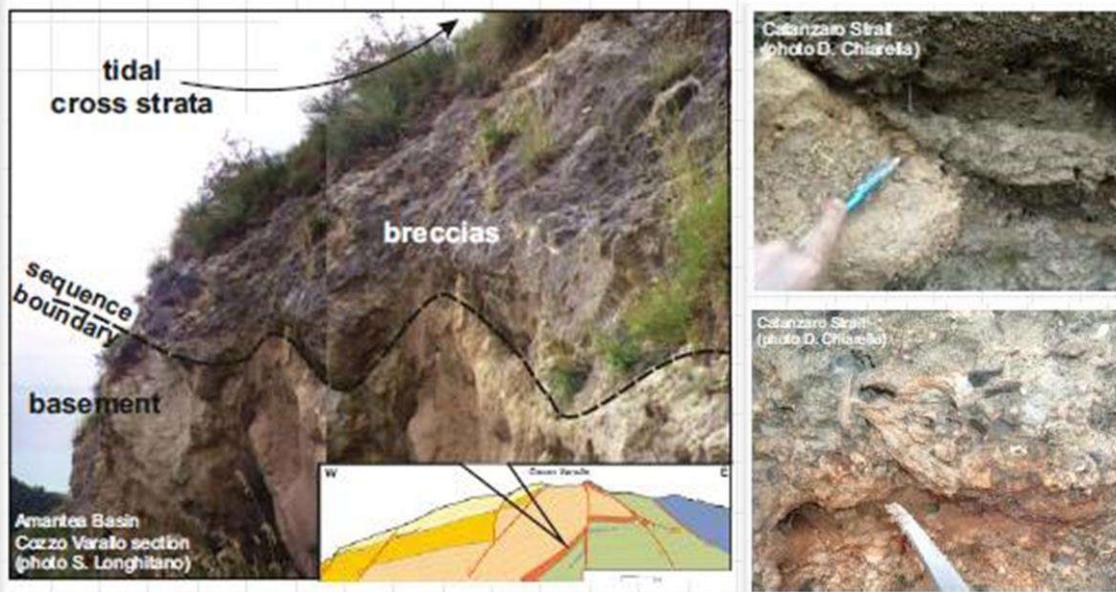
- The Strait-Center Zone



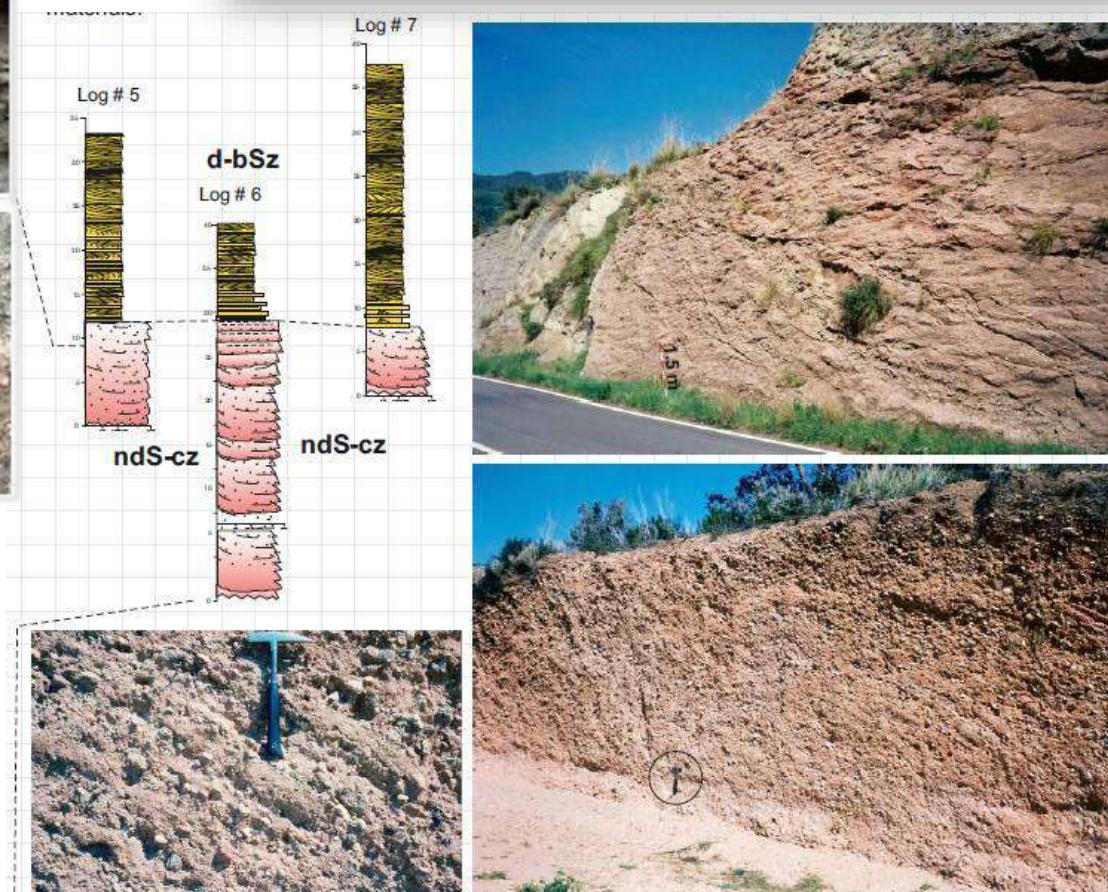
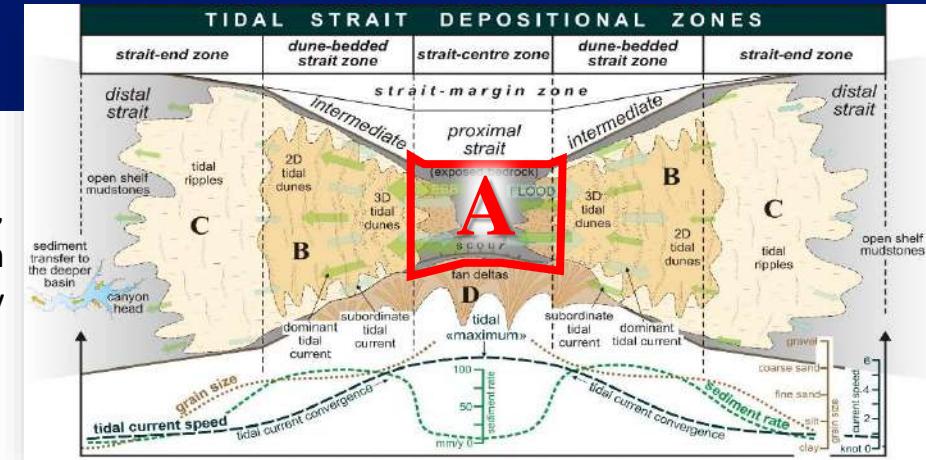
## • The Strait-Center Zone

### 1) GRAVEL/SHELL LAGS: THE STRAIT-CENTER ZONE

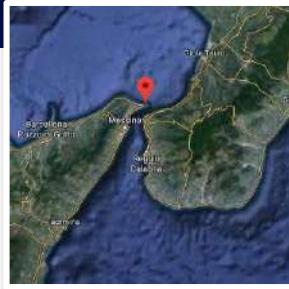
laterally-discontinuous lags 1-2 m thick of massive assemblages of fossil fragments, pebbles and cobbles, immersed in a siliciclastic gravel-size matrix very rich in glaucony. Alternatively, this sector can be highly depositional and represented by coarse-grained sediments derived from the strait margins.



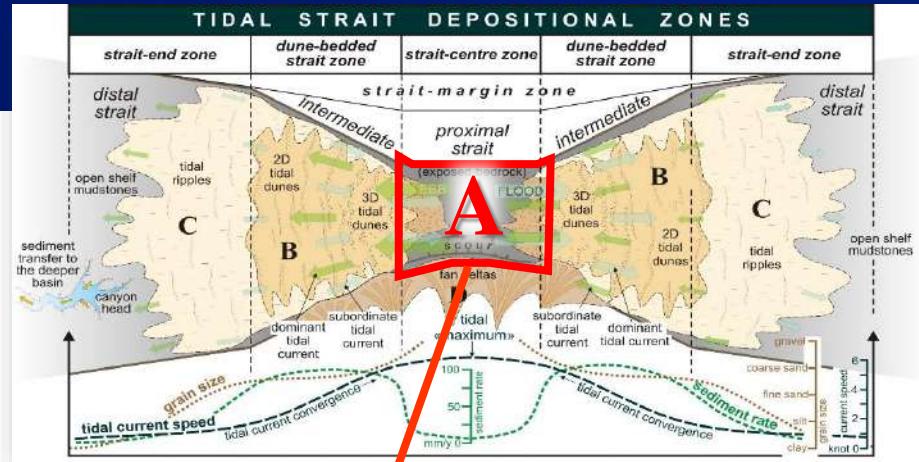
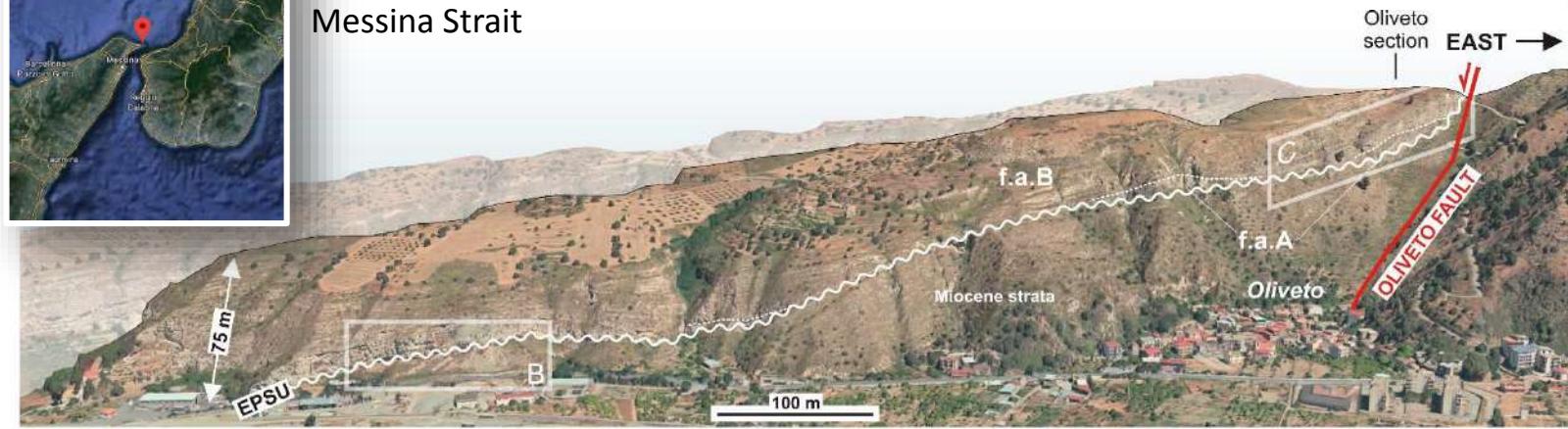
Faunal associations are represented by disarticulated mollusk shells associated with highly-weathered fragments of red algae, *Corallia errinecea* and *Laminaria*, indicating high-energy environments subjected to vigorous currents.sequences.



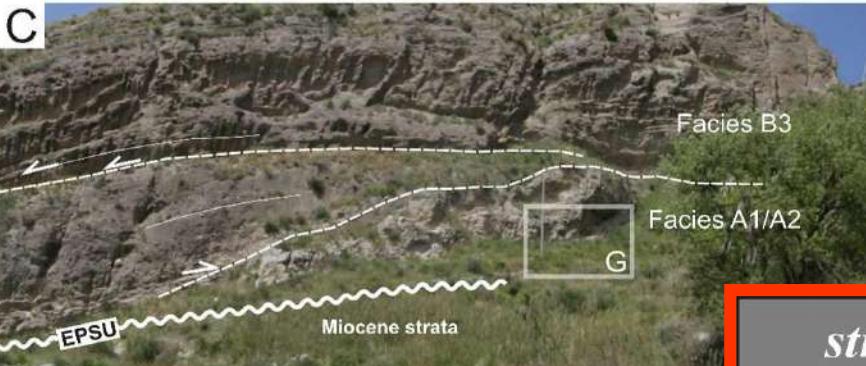
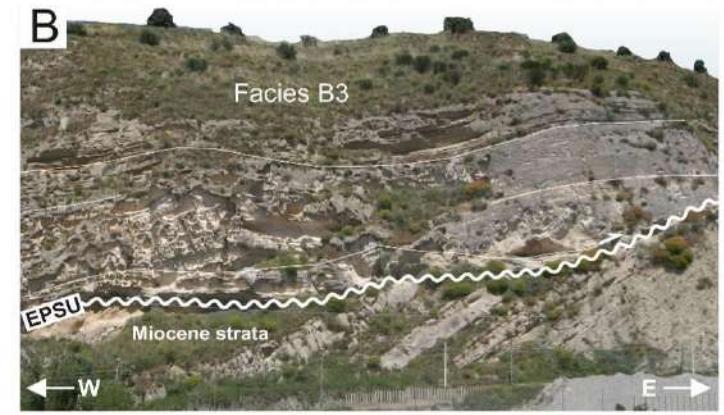
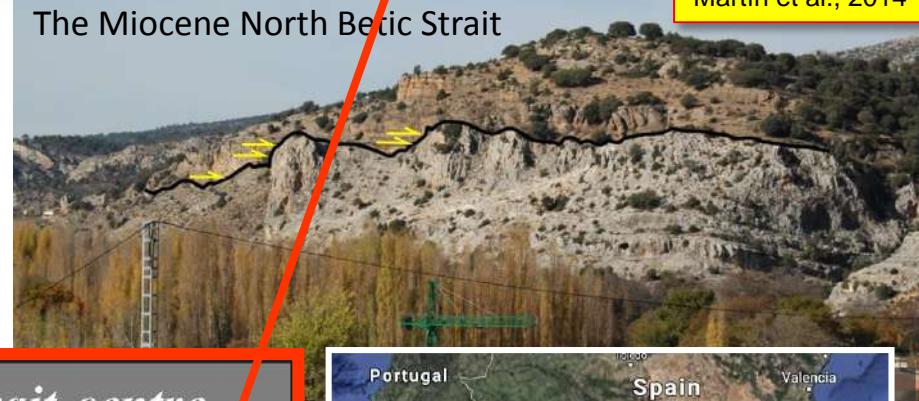
- The Strait-Center Zone



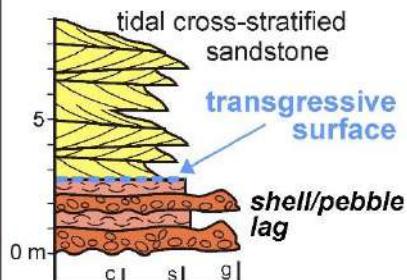
Messina Strait



The Miocene North Betic Strait



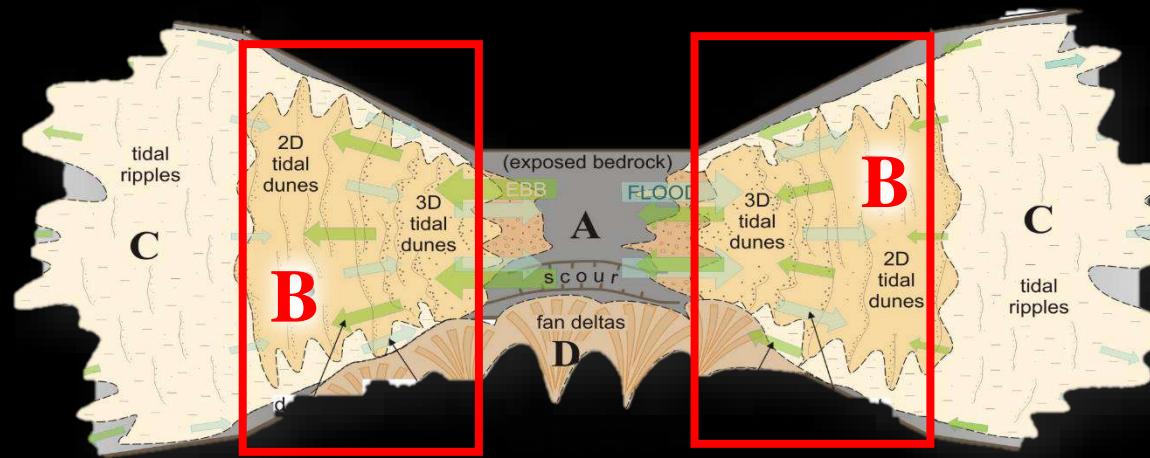
**strait-centre  
(zone A)**



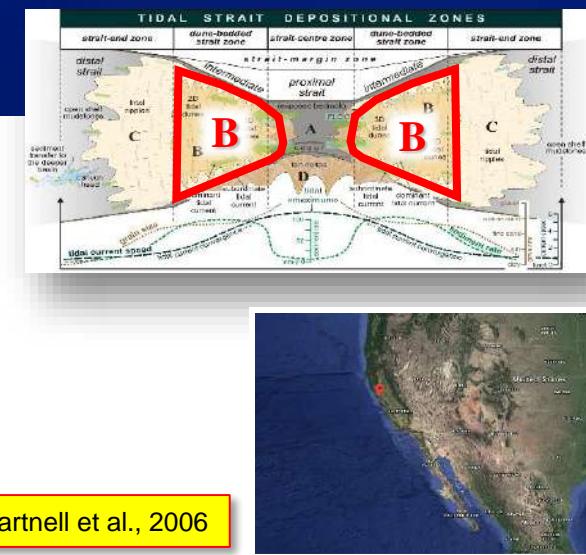
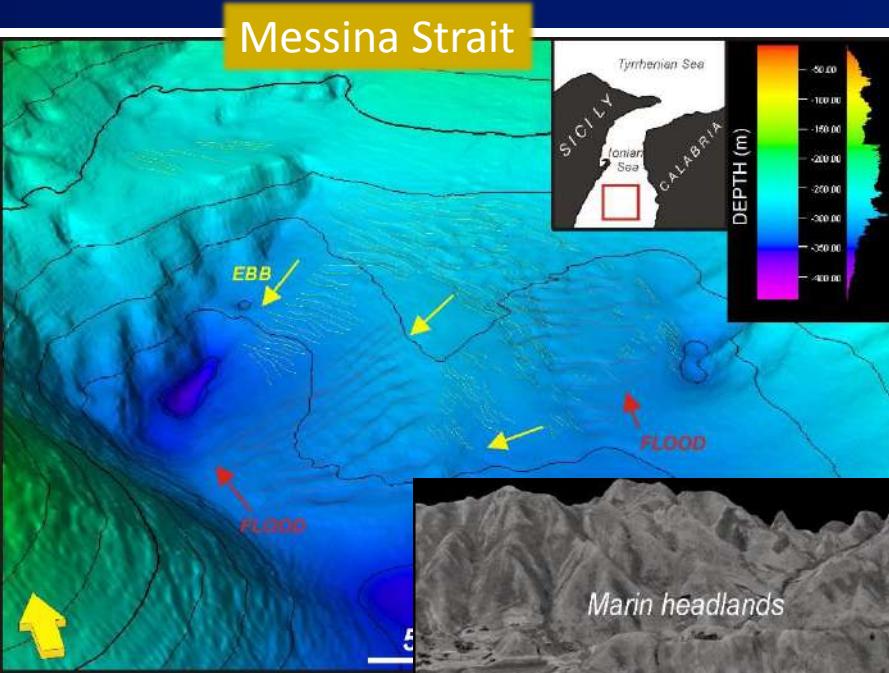
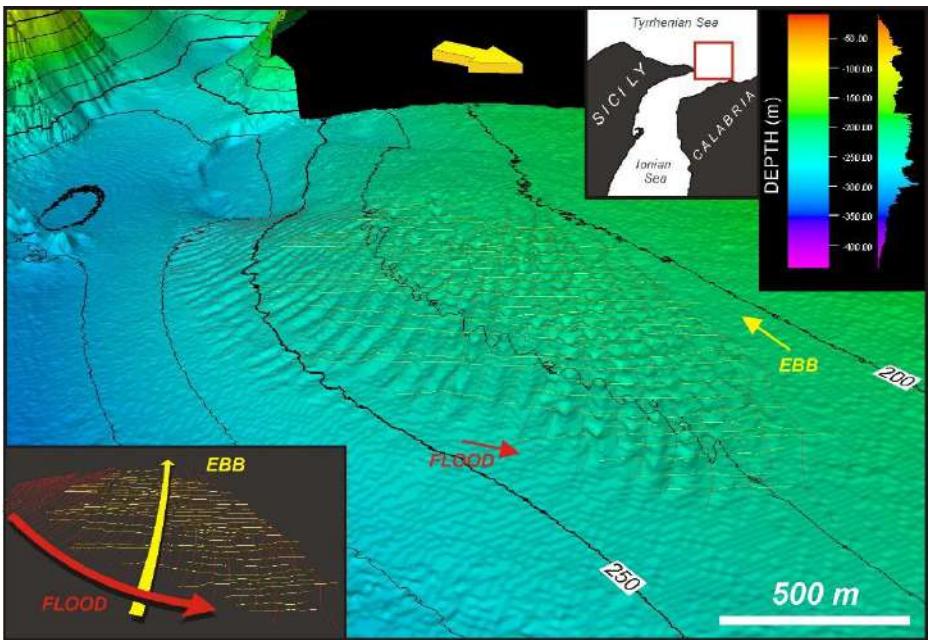
Longhitano, 2018b



# The dune-bedded strait zones



- The Dune-Bedded Strait Zones



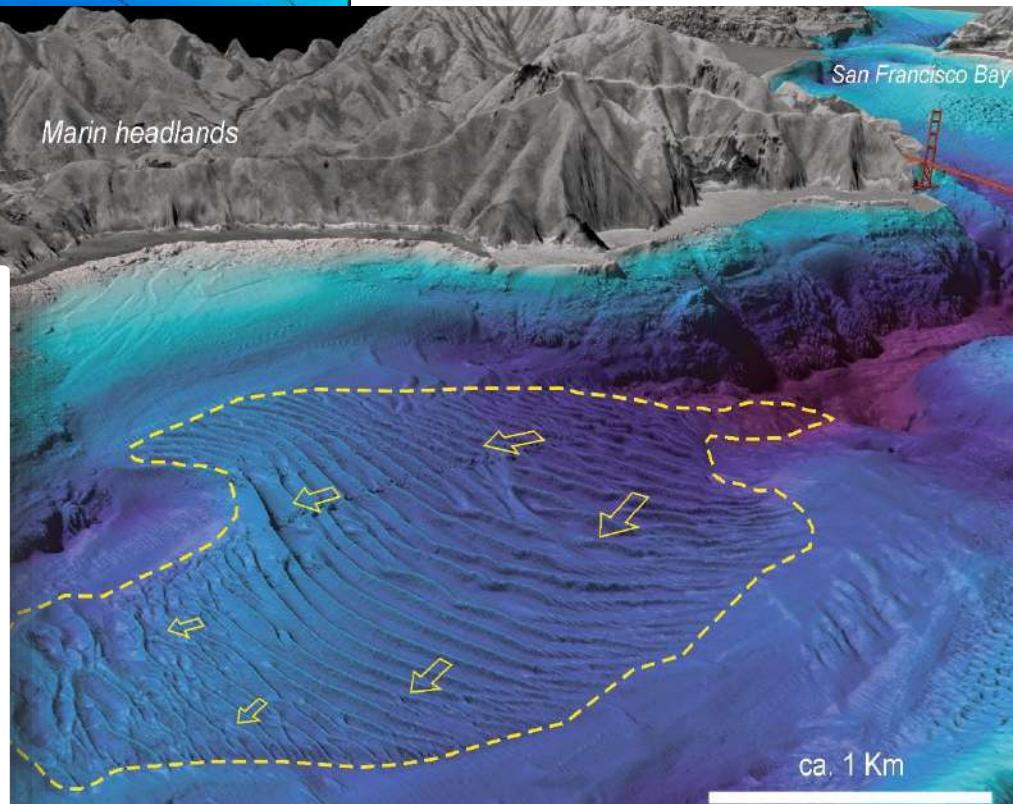
courtesy of A. Colella



courtesy of B. Johnson

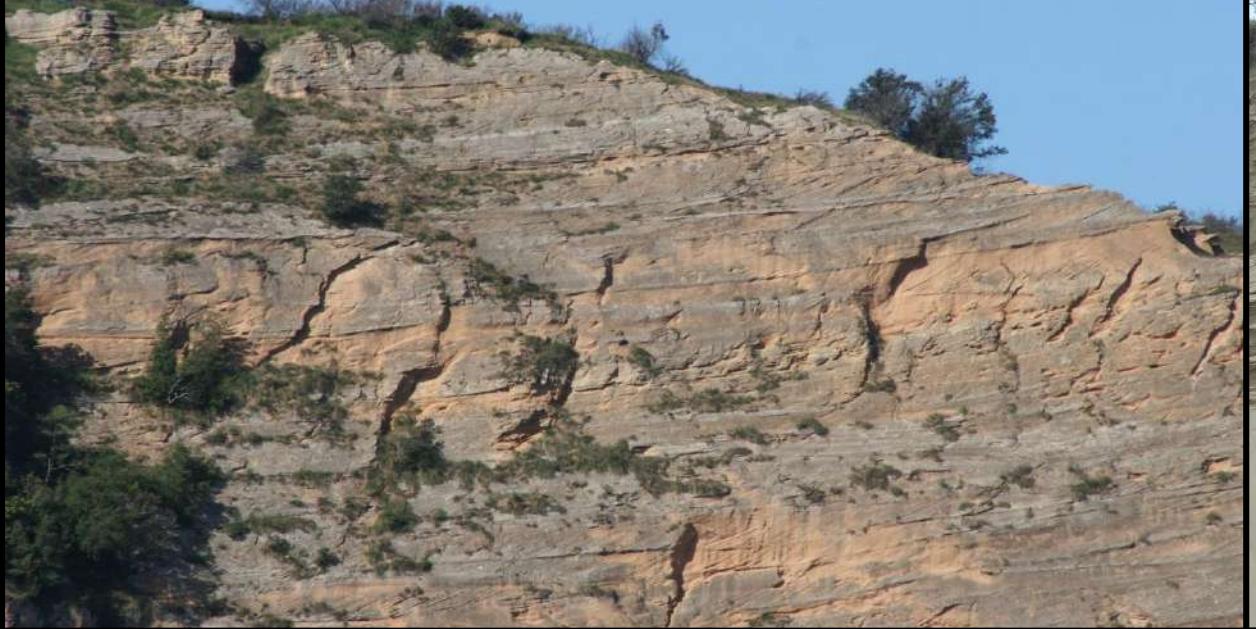


Longhitano, 2018a





AMANTEA STRAIT, TORTONIAN



CATANZARO STRAIT, EARLY PLEISTOCENE



MESSINA STRAIT, EARLY PLEISTOCENE



SIDERNO STRAIT, EARLY PLEISTOCENE

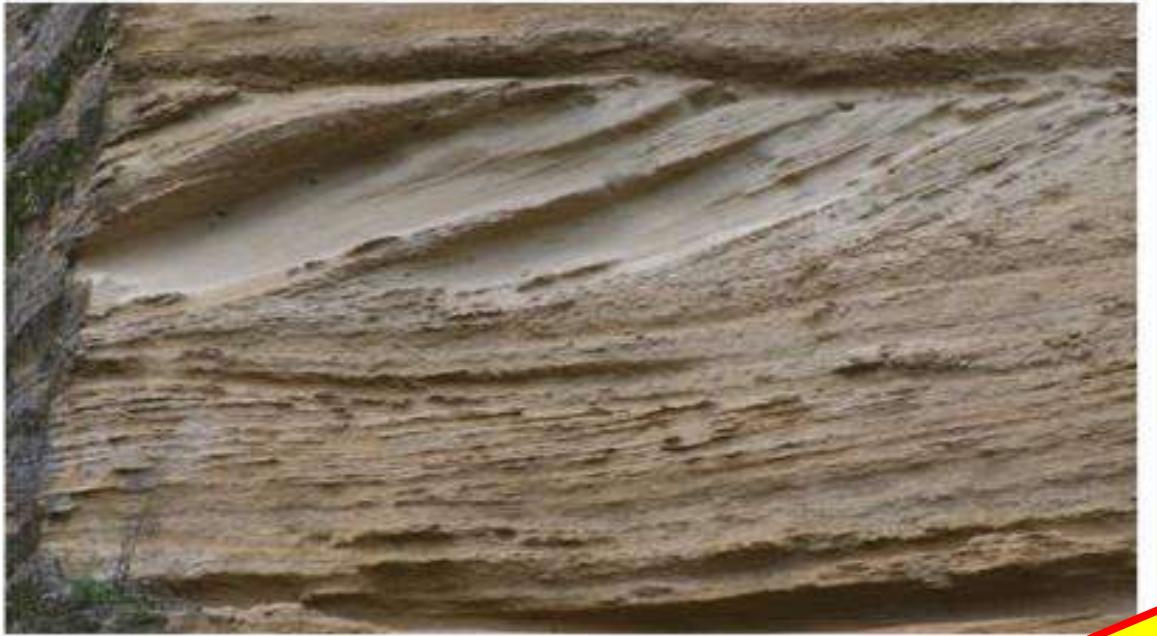


CATANZARO STRAIT, EARLY PLEISTOCENE

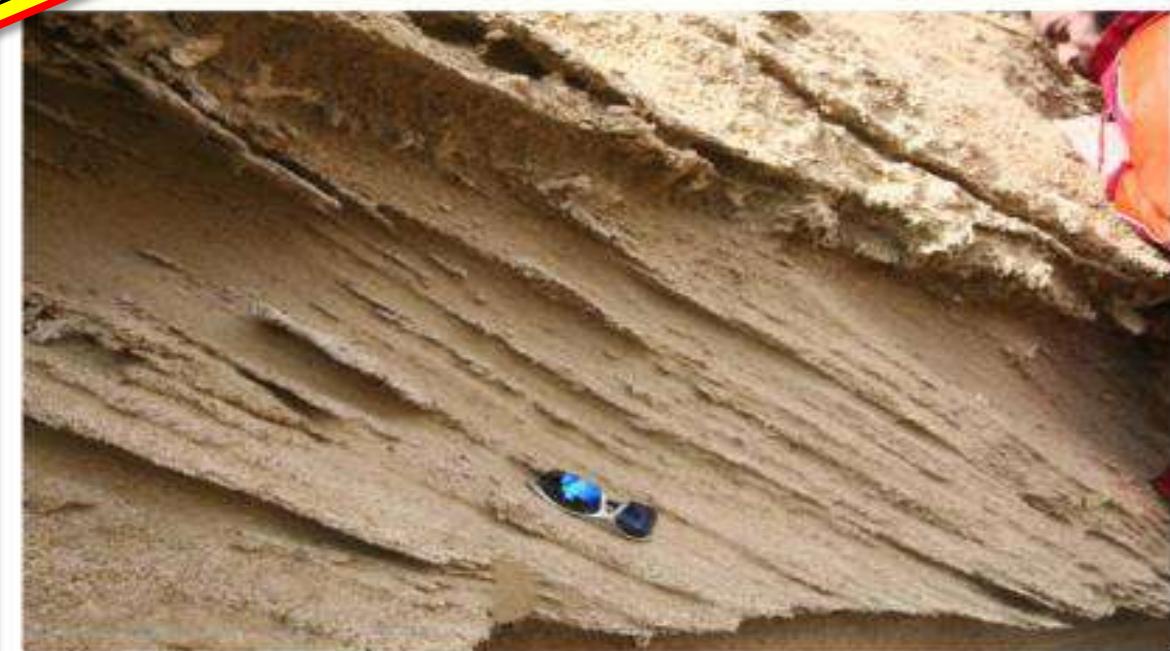


AMANTEA STRAIT, TORTONIAN





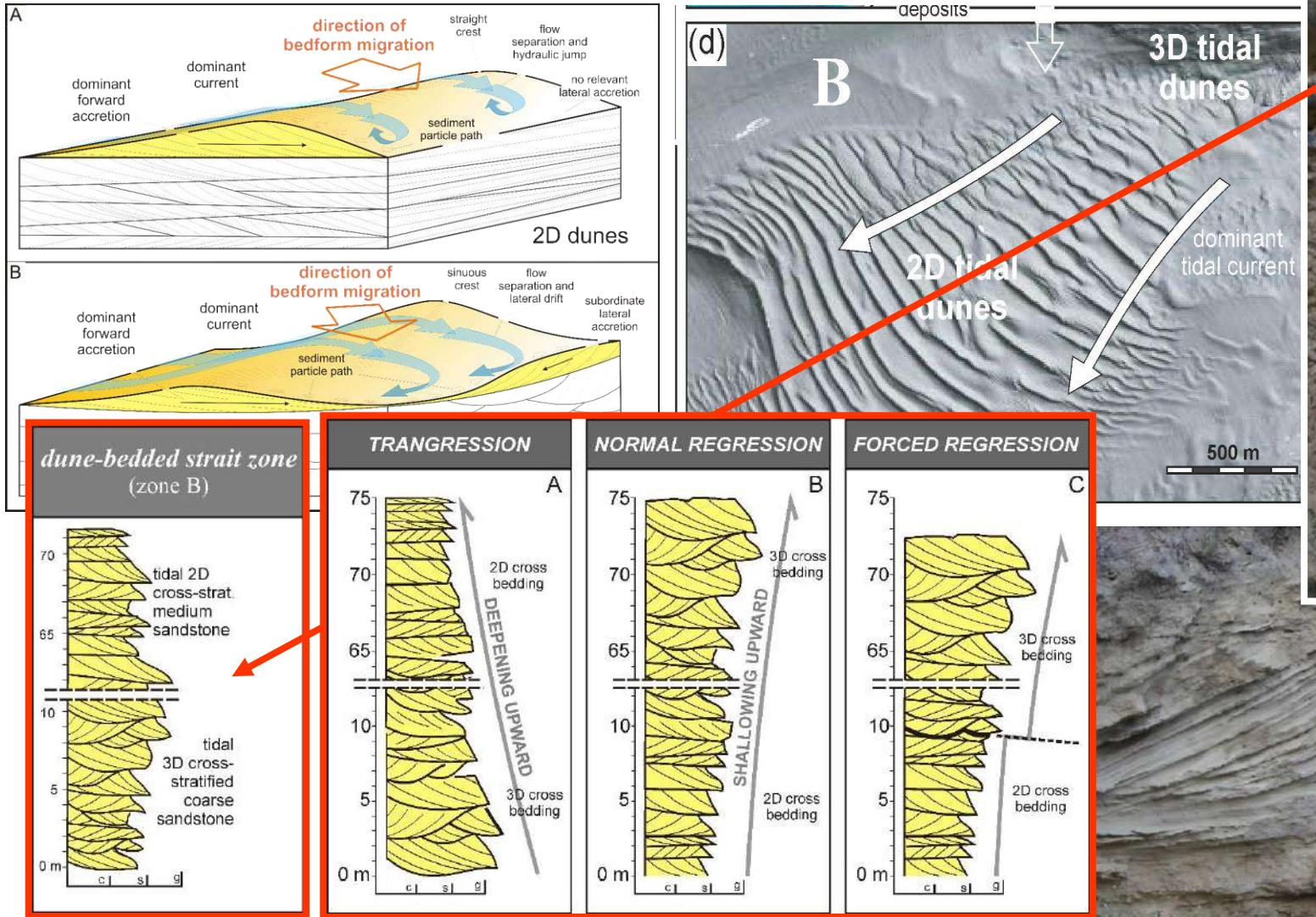
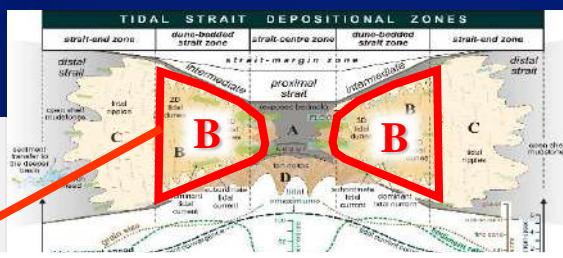
A VARIETY OF INTERNAL SEDIMENTARY STRUCTURES  
CONCUR IN SUPPORTING THE INTERPRETATION OF  
TIDE-DOMINATED DEPOSITS



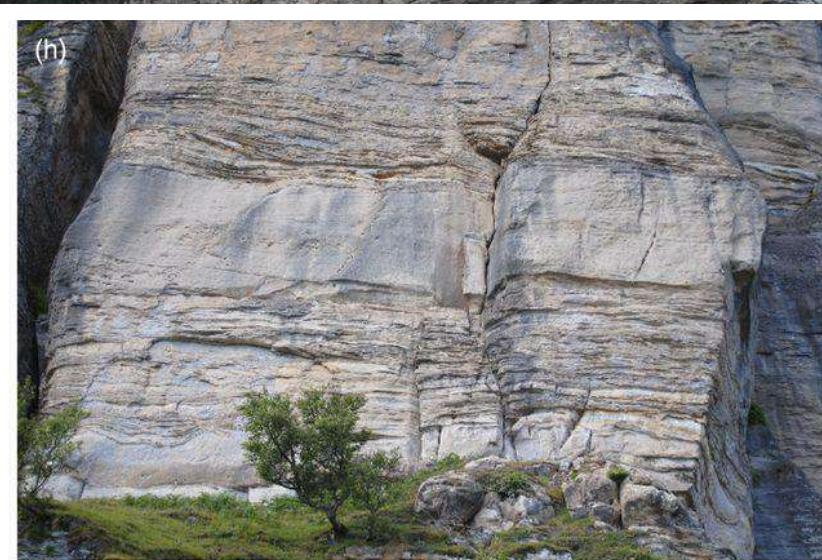
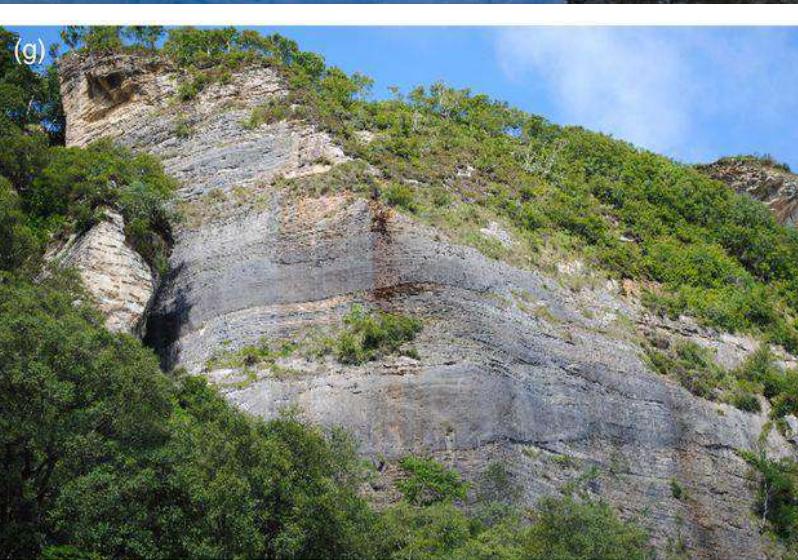
## • The Dune-Bedded Strait Zones

### 2D/3D TIDAL CROSS STRATA

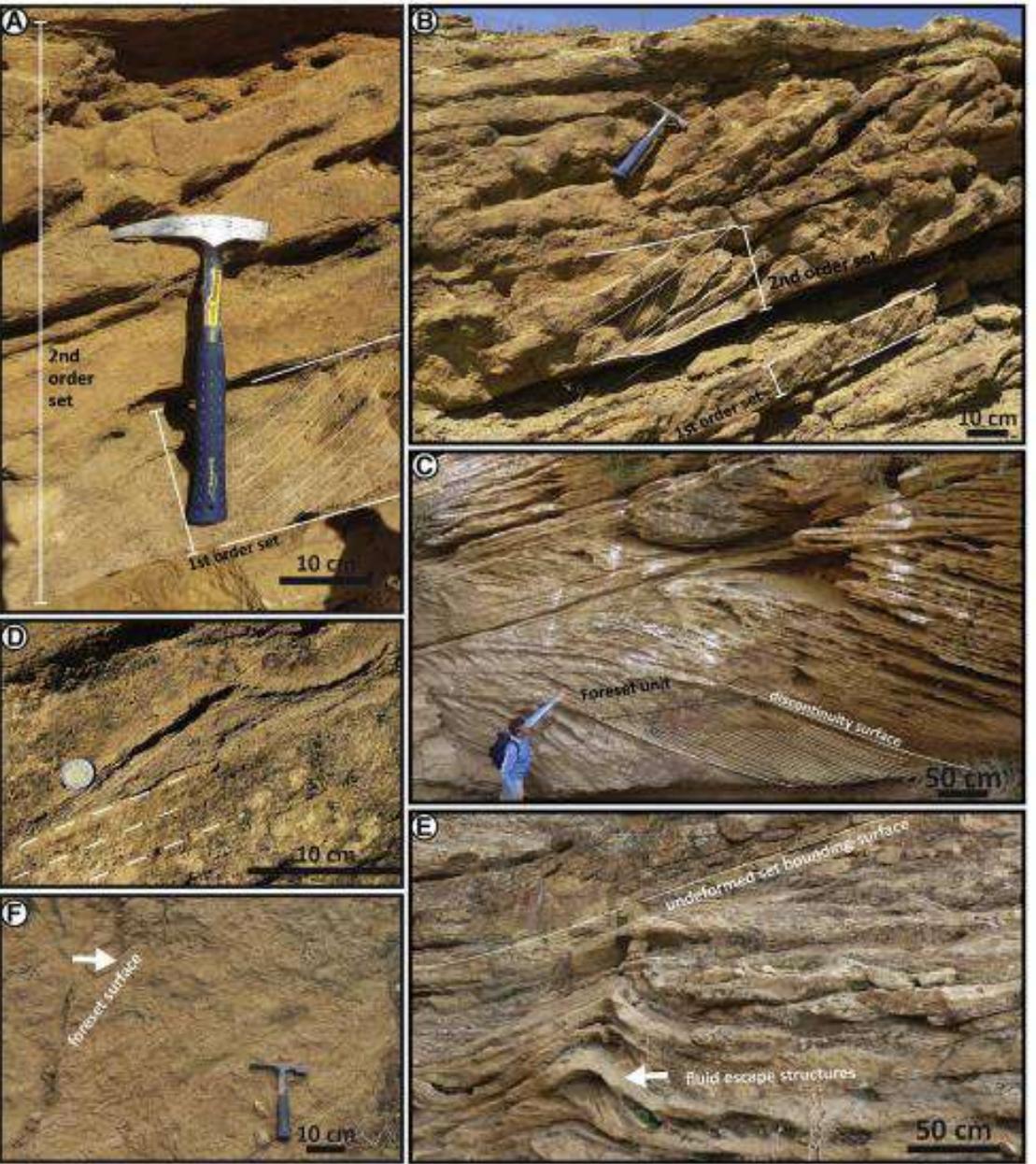
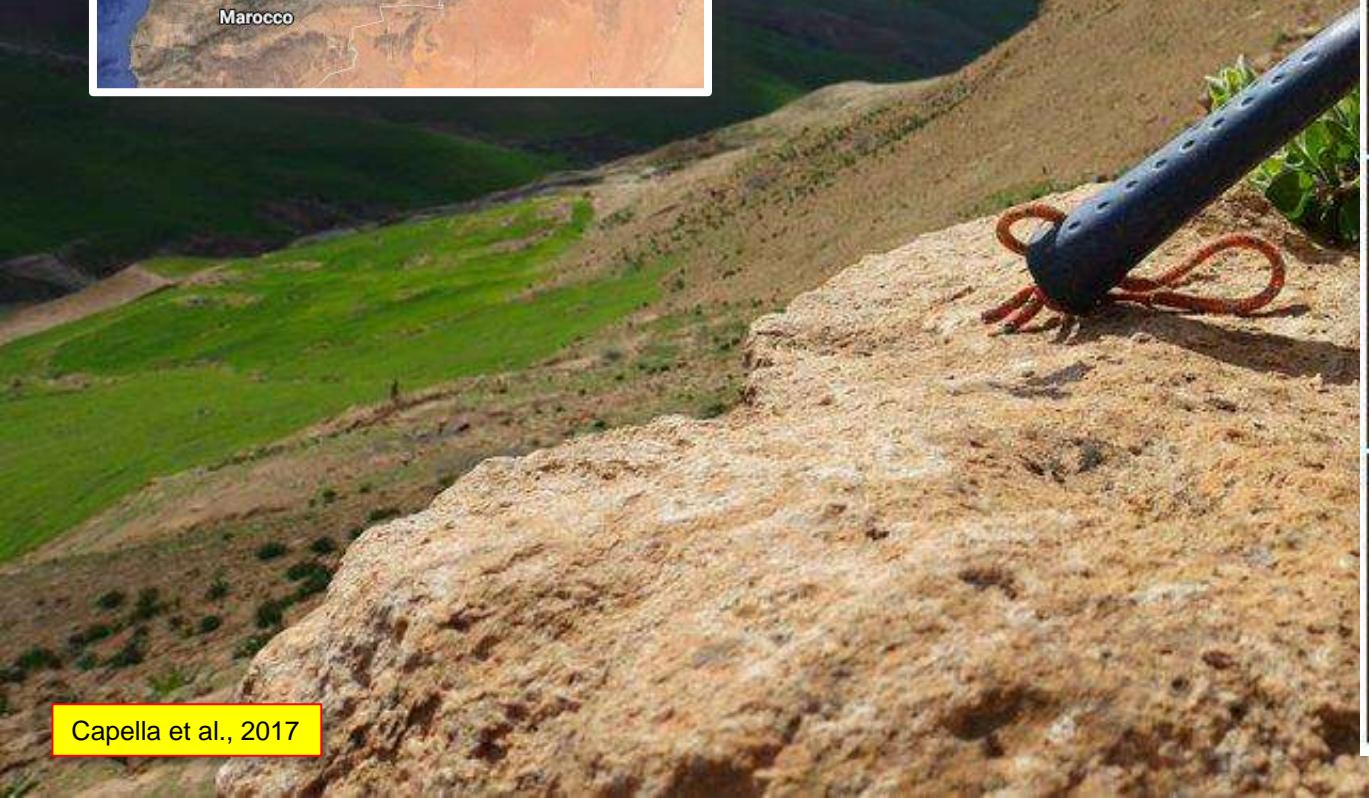
Cross strata sets can be differentiated into **trough** and **planar** architectures, indicating sinuous- and straight-crested bedforms (i.e., three- vs. two-dimensional) as they are believed to represent **energetic transitional flow stages**.



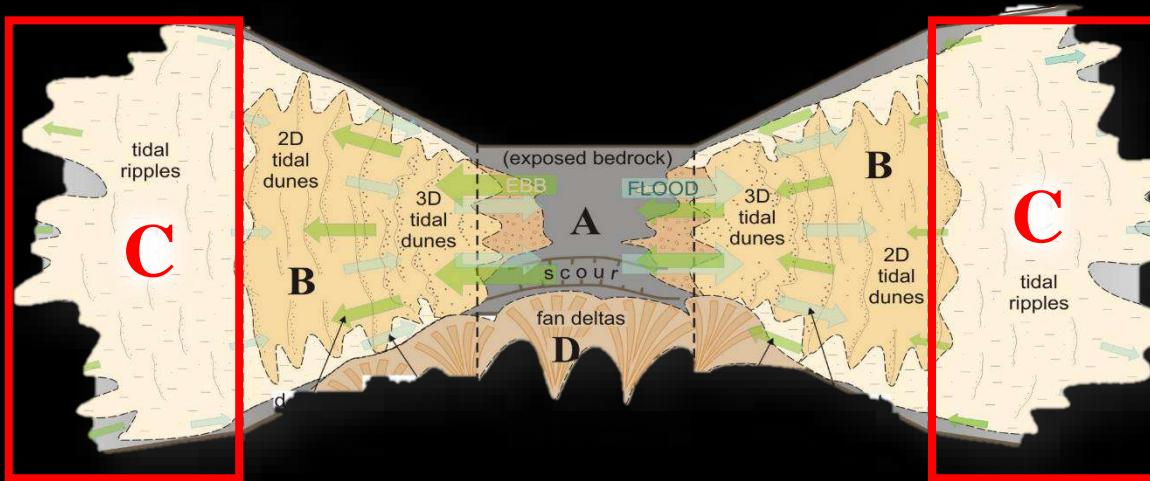
# Middle Jurassic of the Isles of Skye and Raasay, NW Scotland



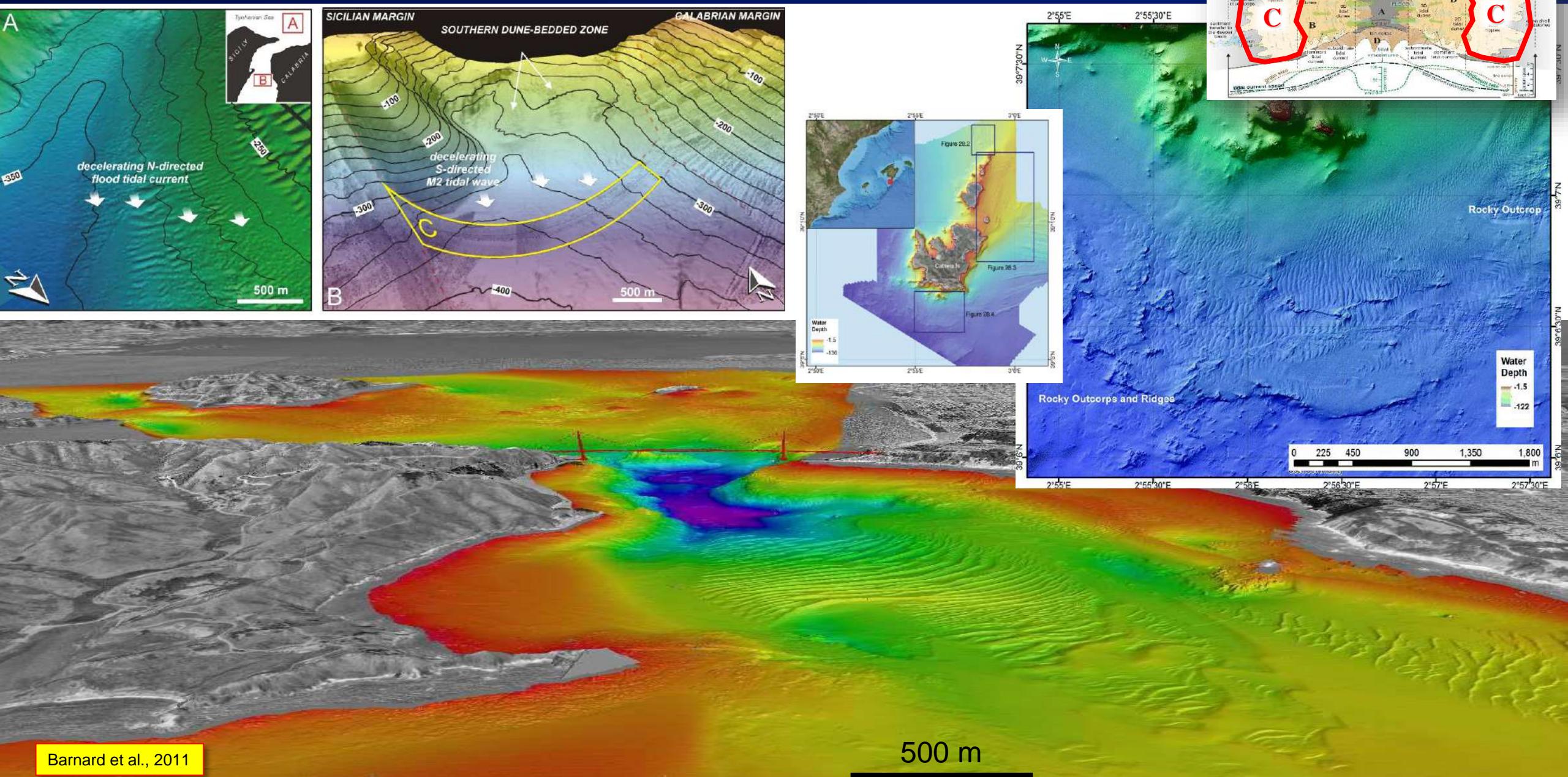
# Upper Miocene of the Rifian Corridor, northern Morocco



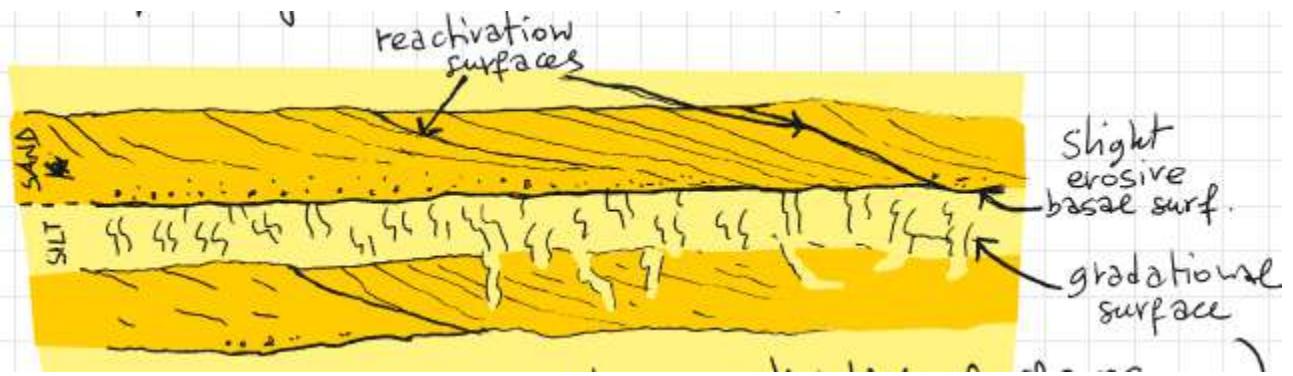
# The strait-end zones



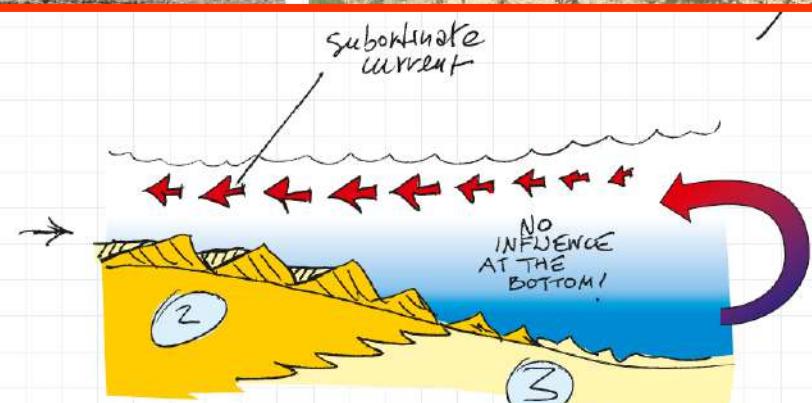
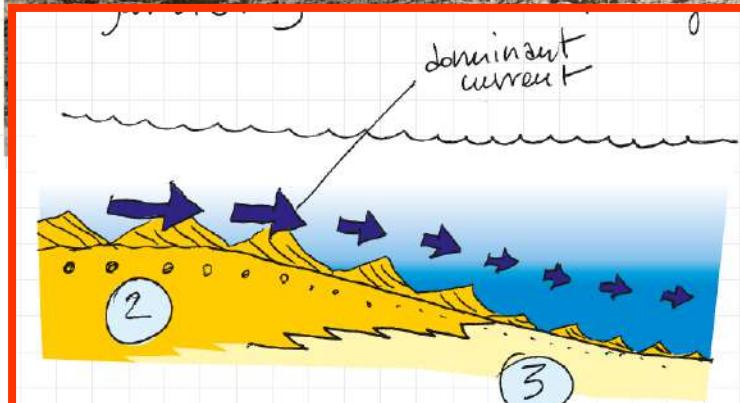
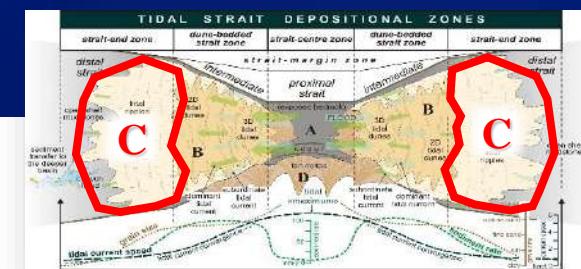
- The Strait-End Zones



- The Strait-End Zones



*Skolithos*,  
*Thalassinoides*,  
*Psilonichnus* and  
*Arenicolites* trace  
fossils

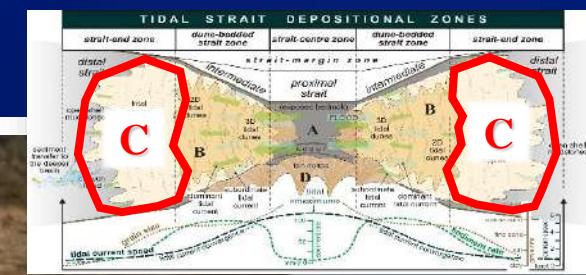




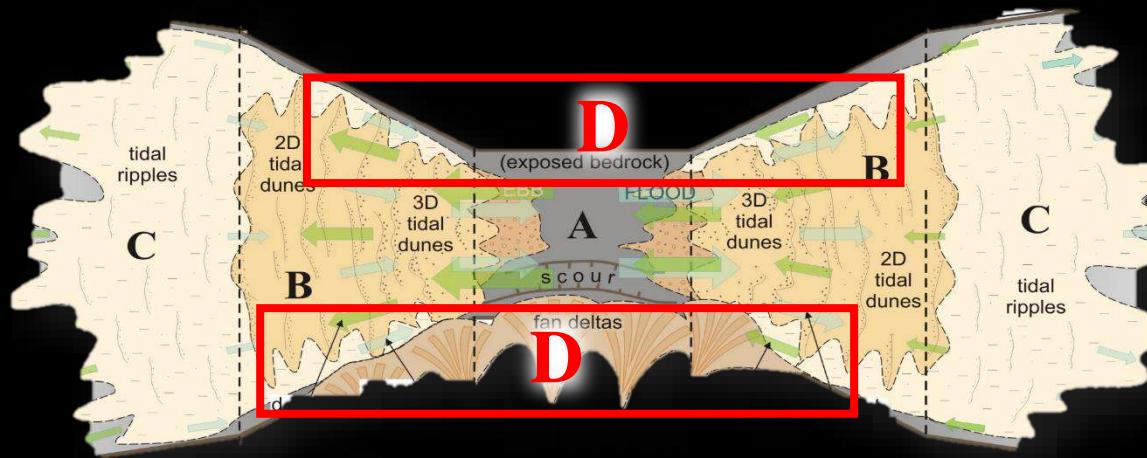
CATANZARO STRAIT, EARLY PLEISTOCENE

- The Strait-End Zones

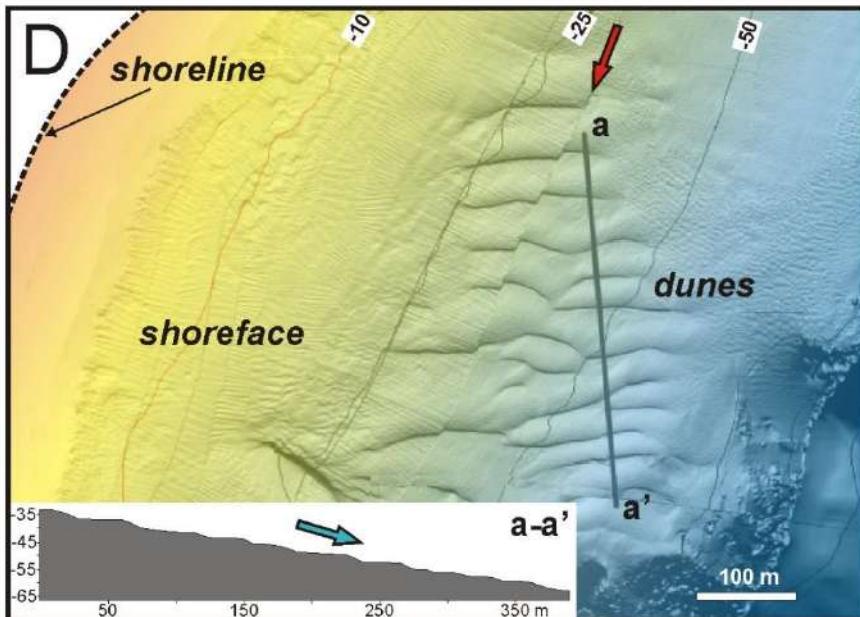
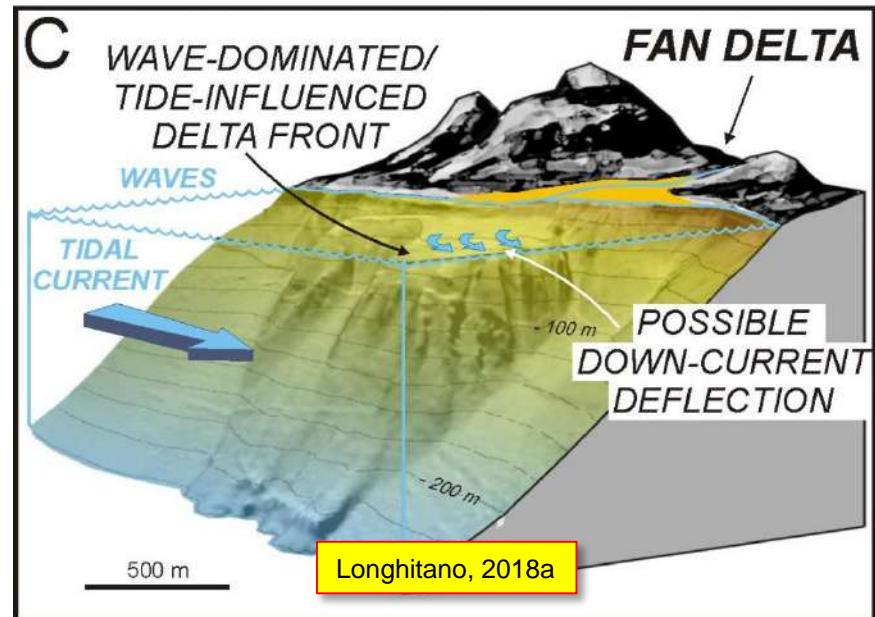
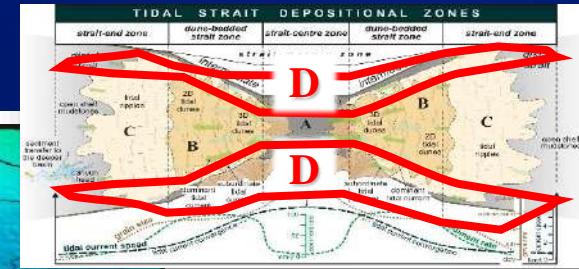
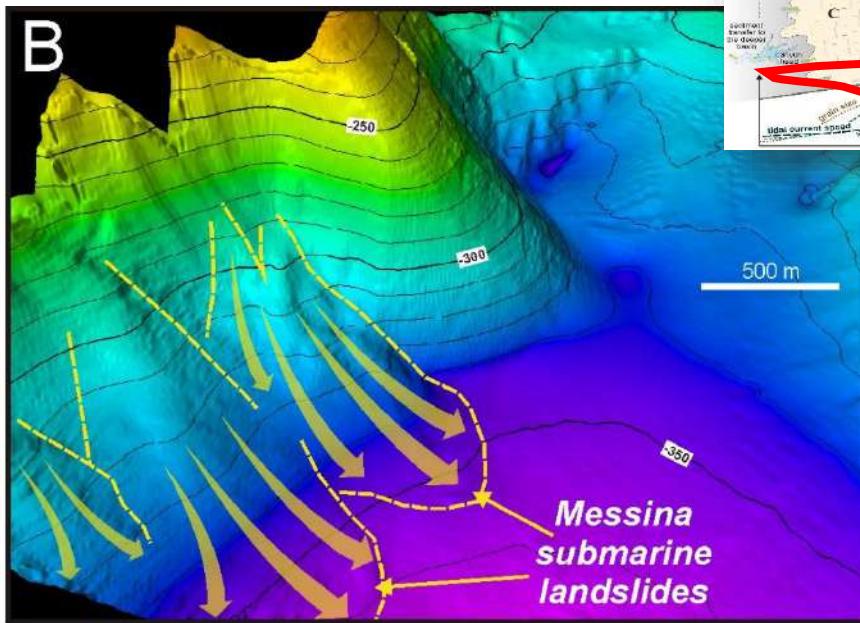
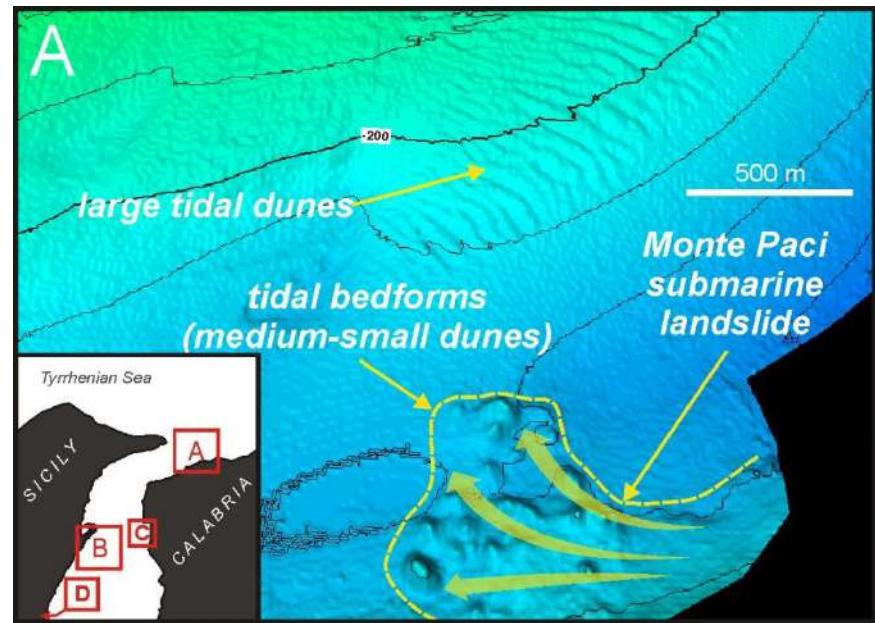
MESSINA STRAIT, EARLY PLEISTOCENE



# The strait-margin zones



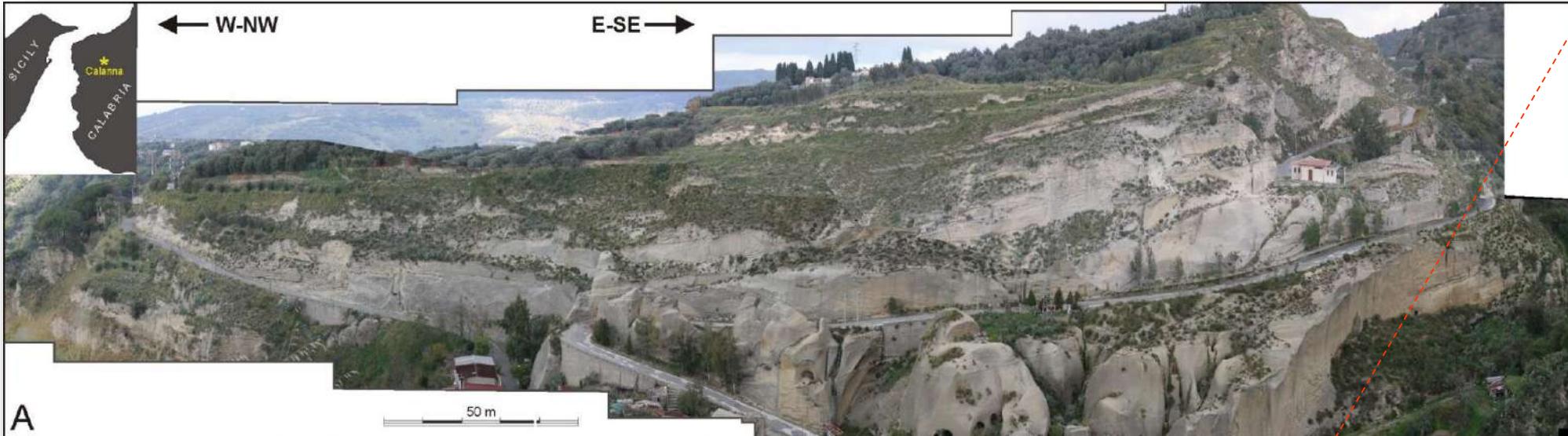
- The Strait-Margin Zones



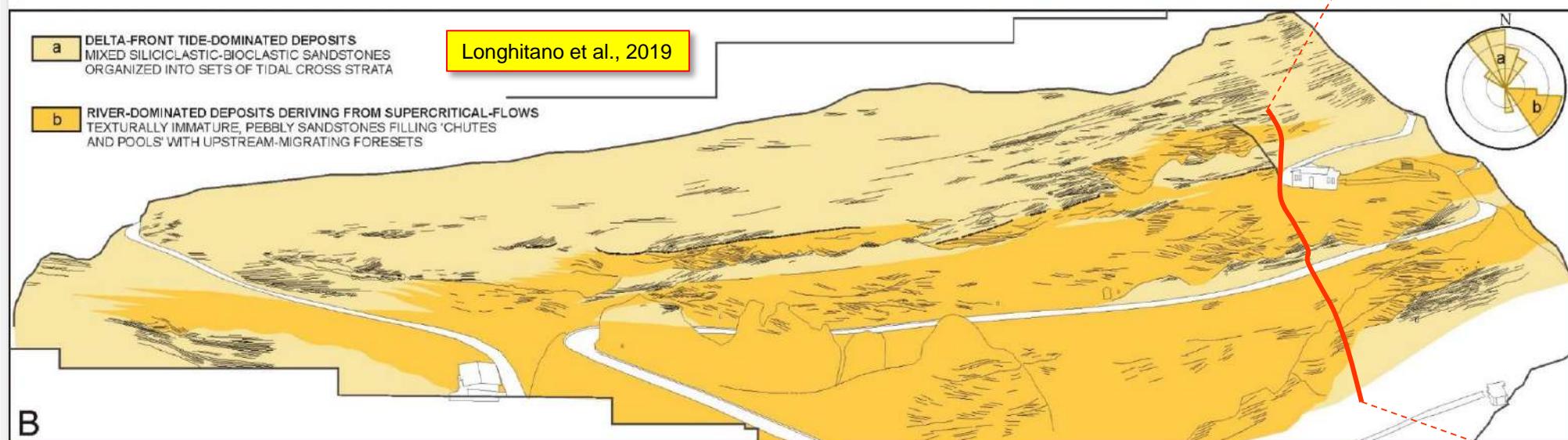
- The Strait-Margin Zones

Barrier et al., 1987

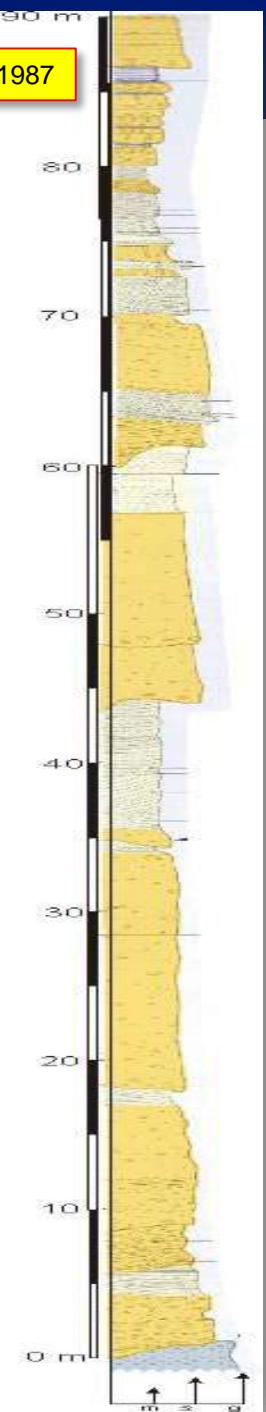
MESSINA STRAIT, EARLY PLEISTOCENE



A



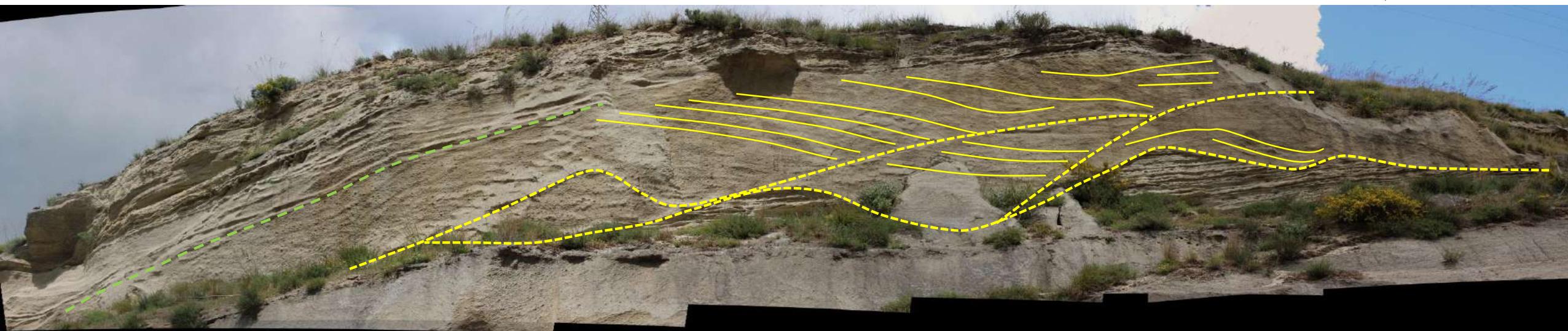
B



**Two main stratigraphic surfaces:**

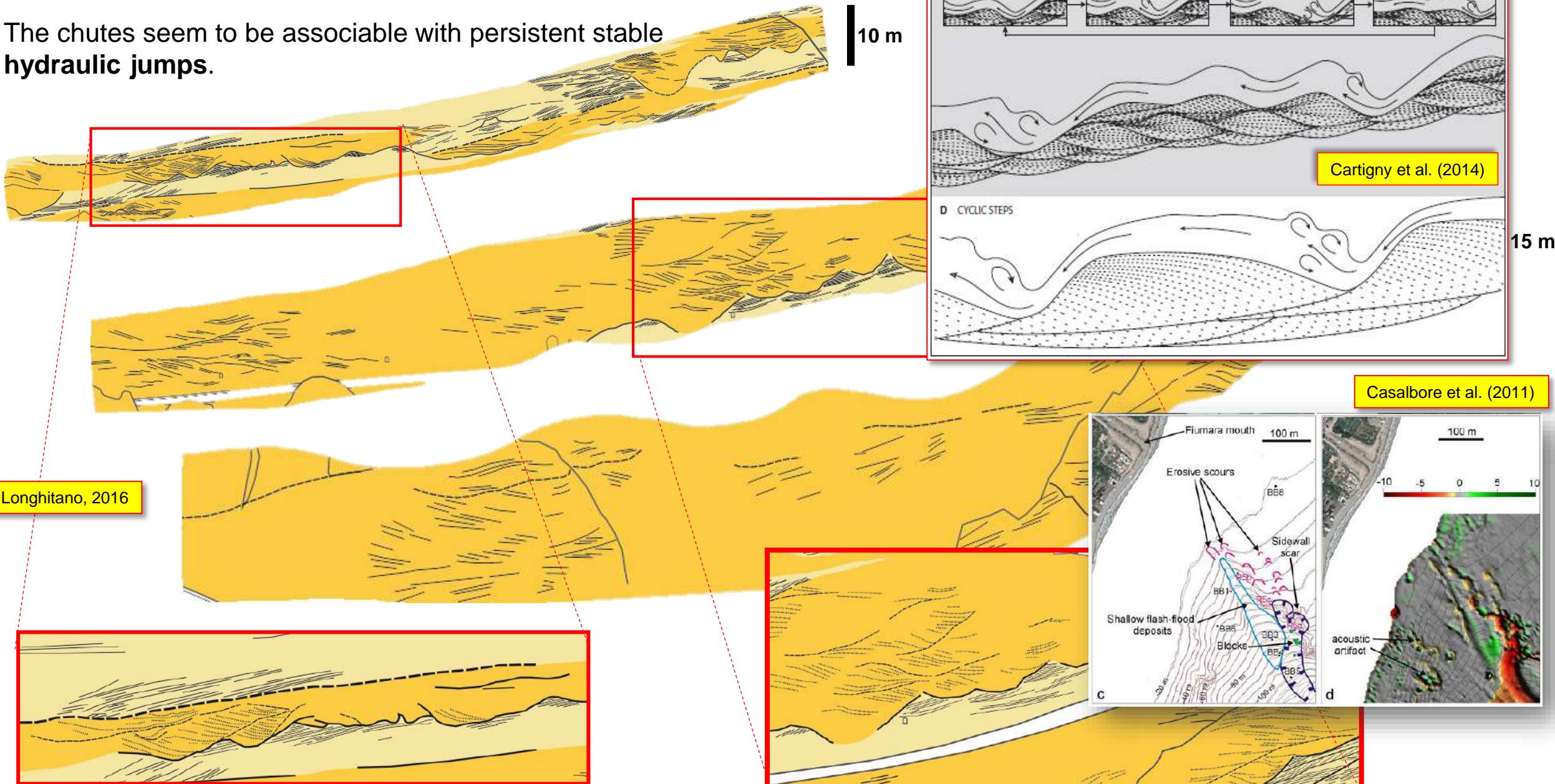
- a. erosional (i.e., scour and fill, some m deep, multiple scours, irregularly spaced)
- b. depositional (i.e., tidal reworking, unidirectional foresets)

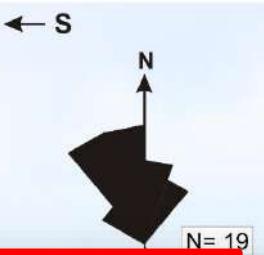
MESSINA STRAIT, EARLY PLEISTOCENE



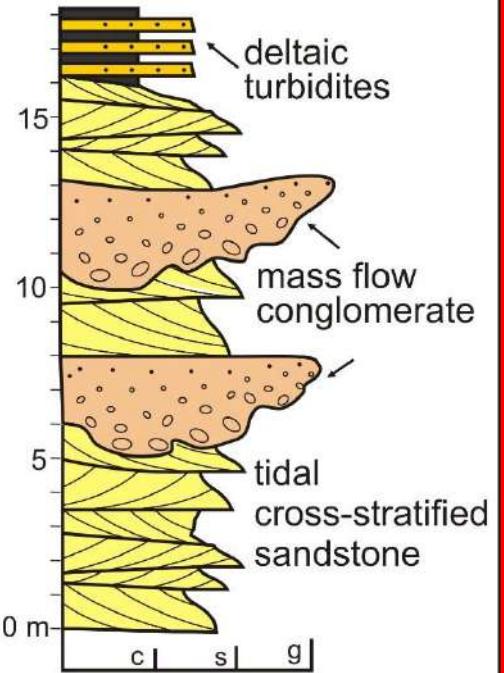
Upstream migrating chutes that end downstream in a series of unstable antidunes. These features indicate repeated hydraulic jumps and surges suggesting **chutes-and-pools** structures.

The chutes seem to be associated with persistent stable **hydraulic jumps**.





## *strait-margin* (zone D)



Longhitano et al., 2012b



# SUMMARIZING ALL WE TODAY KNOW ABOUT THE SEDIMENTARY DYNAMIC OF TIDAL STRAITS ...

ZONE of CURRENT SLACK and INVERSION	ZONE of CURRENT ACCELERATION / DECELERATION	ZONE of CURRENT AMPLIFICATION	ZONE of CURRENT ACCELERATION / DECELERATION	ZONE of CURRENT SLACK and INVERSION					
TIDAL POWER max									
TIDAL PHASE flood ebb	symmetric	slightly symmetric	strongly symmetric	symmetric					
TIDAL PHASE symmetric	slightly symmetric	strongly symmetric	symmetric	symmetric					
SEDIMENTARY STRUCTURES / ARCHITECTURES	CLASTIC CARBONATE WEDGES	ALTERNANCE of CHANNEL-FILLS and MEDIUM-SCALE CROSS BEDDING	CANYON-FILL FINING-UPWARD SEQUENCE	NEPTUNIAN DIKES SCOURS GRAVEL CHANNEL-FILLS	GRAVEL LENSES / PATCHES	UPPER-STAGE PLANE BEDS and HERRINGBONE CROSS BEDDING DOMINANTLY BI-DIRECTIONAL	LARGE-SCALE CROSS BEDDING DOMINANTLY UNI-DIRECTIONAL	MEDIUM-SCALE CROSS BEDDING EQUALLY UNI- and BI-DIRECTIONAL	SMALL-SCALE CROSS BEDDING MODERATELY BI-DIRECTIONAL
GRAIN-SIZE / LITHOLOGY	GRAVEL and COARSE SAND with FINE-GRAINED INTERCALATIONS, in places		GRAVEL, GRANULE and VERY COARSE SHELLY SAND		GRAVEL, GRANULE and VERY COARSE SHELLY SAND	COARSE SAND	COARSE-MEDIUM SHELLY SAND	HETEROLITHIC MEDIUM-FINE SAND and MUD	SANDY INTERCALATIONS
DEPOSITIONAL ZONES	STRAIT-MARGIN ZONE		STRAIT-CENTRE ZONE		DUNE-BEDDED STRAIT ZONE		STRAIT-END ZONE		

- Concluding remarks

STRAITS (TIDAL AND NON-TIDAL) ARE **INDIVIDUAL DEPOSITIONAL SYSTEMS** STILL POORLY KNOWN.  
ALTHOUGH WE KNOW EVEN MORE ON SUCH SYSTEMS, THEY STILL SUFFER OF A **GENERAL ABSENCE OF UNIVERSALLY APPLICABLE MODELS**, NECESSARY TO INTERPRET MODERN AND ANCIENT EXAMPLES.

HOWEVER ...

- WE KNOW THEIR HYDRODYNAMICS AND TIDAL MOTION IN A QUITE ACCURATE WAY AND PROBABILISTIC NUMERICAL MODELS SEEM TO WORK WELL.
- STRAIT BEDFORMS INDICATE *CONTINUUM* OF TYPES AND GEOMETRIES/SIZES, BUT THERE ARE STILL POORLY KNOWN ELEMENTS, PROBABLY REFERABLE TO THE RECENT GEOLOGICAL HISTORIES OF OUR MODERN STRAITS.
- THE POTENTIAL OF THE STUDY OF ANCIENT ANALOGUES IS HUGE: STRAIT-FILL SUCCESSIONS MAY REVEAL, INDEED, PROCESSES NOT OBSERVABLE AT HUMAN TIME-SCALE AND ALLOW INTEGRATING/IMPROVING KNOWLEDGE OF SUCH COMPLEX SYSTEMS.



Courtesy of  
Silvia Messina  
UNIBAS Erasmus student at Université de la Normandie,  
Caen, France

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